

# The Age of the Earth — NOT!

~~Why is it so difficult to determine the age of the Earth?~~

## Absolute versus relative ages

Geologists have long been able to determine relative ages from cross-cutting relationships of rock units in outcrop.

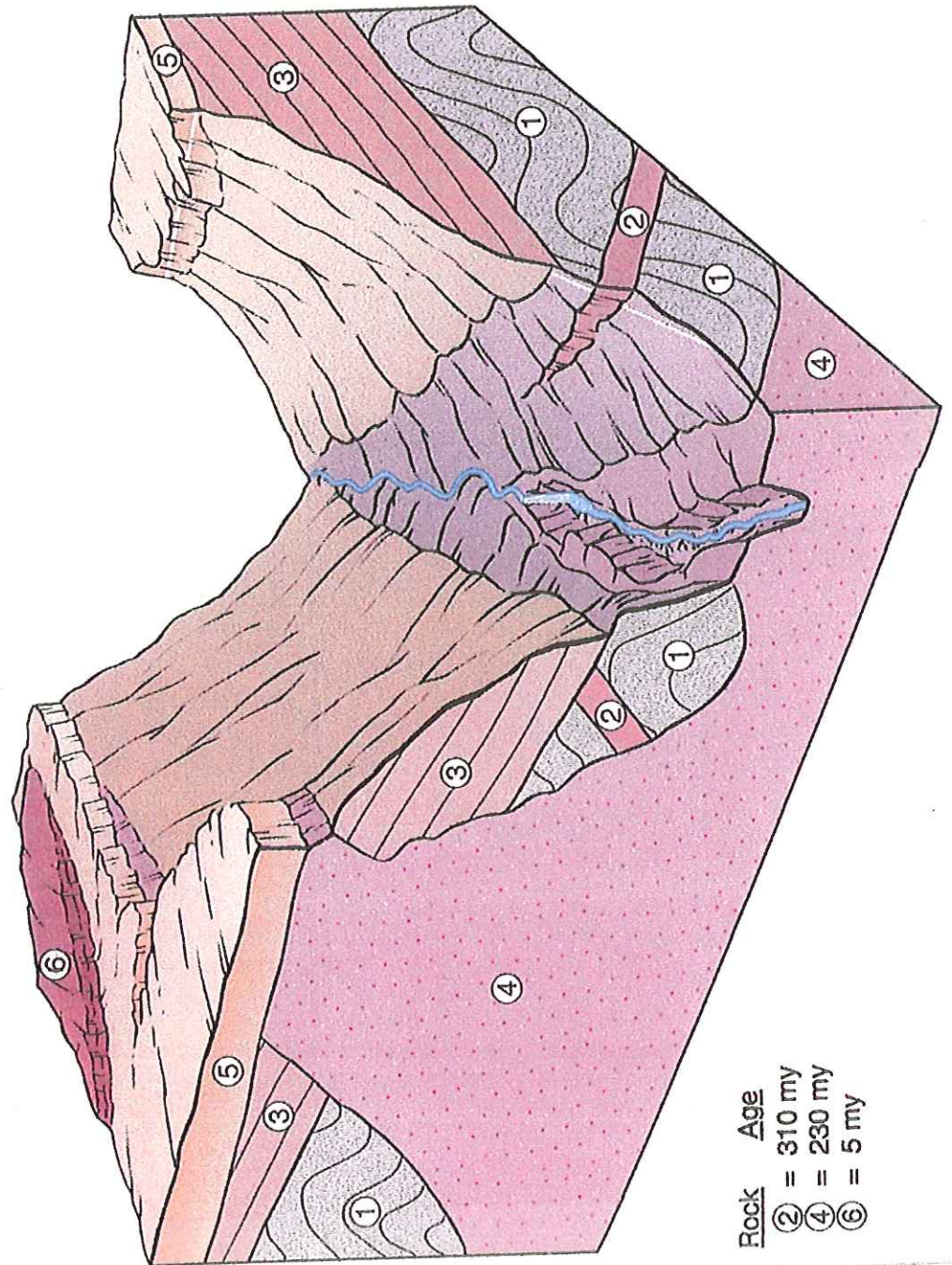
Cartoon example Fig. 8.4 — discuss briefly  
In practice, it requires careful observation to recognize such relations in the field

Absolute ages are more difficult — but there have been many attempts to answer the fundamental question — how old is the  $\oplus$ ?

↳ also T&R perspective 7.2

Figure 8.1 summarizes some of these ~~ideas~~ since the Renaissance — and with an emphasis on European ideas

Biblical chronologies exemplified by Archbishop James Ussher, ~~of~~ Ireland 1581 — 1656



**FIGURE 7.19**

Radiometric dates obtained on igneous rocks can be fitted into the geologic time scale based on sedimentary rocks. This is done by bracketing the sedimentary rocks with the igneous rocks. In some instances radiometric dates can be assigned directly to the sedimentary rocks. See text for discussion.



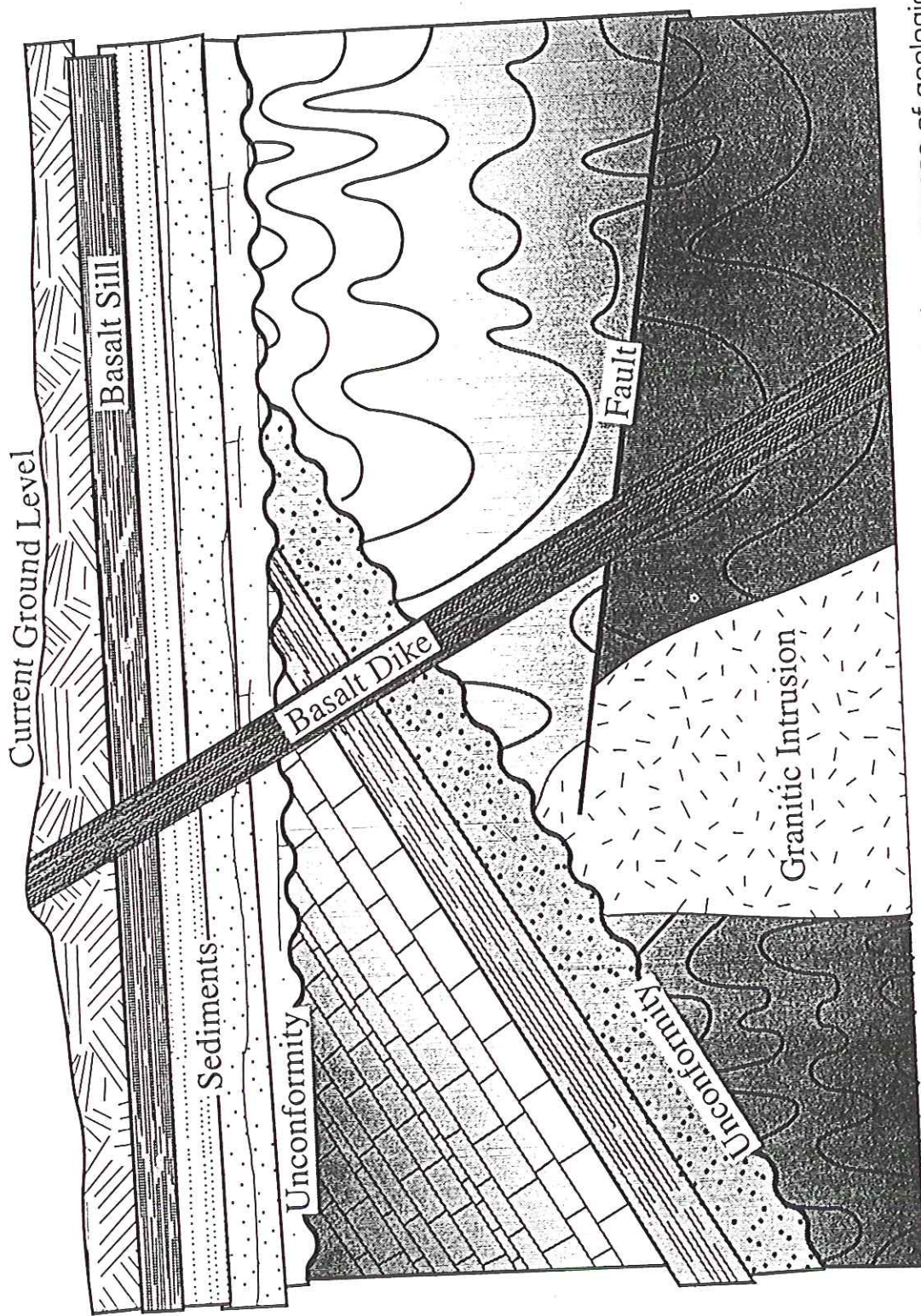
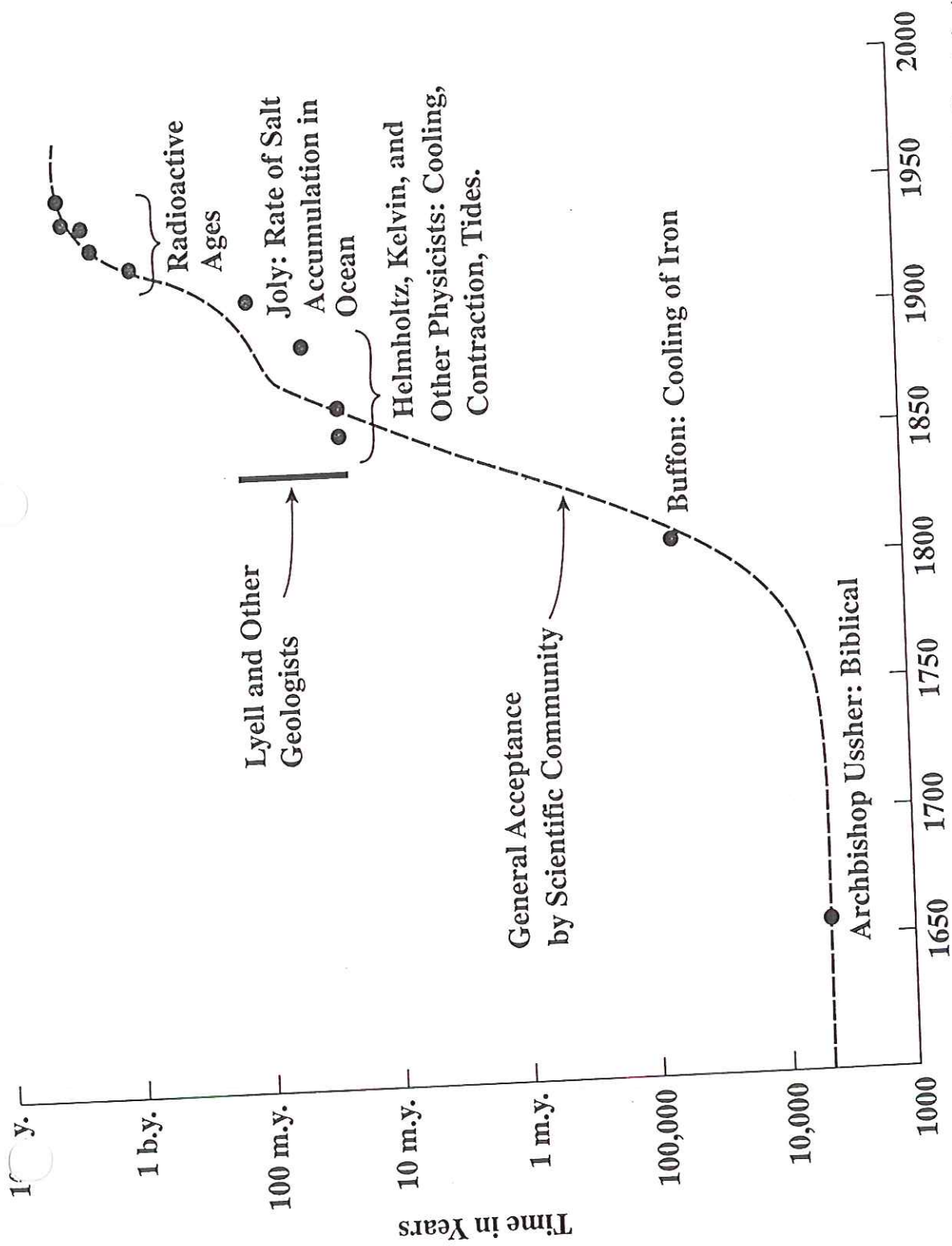


Figure 8.4. Idealized section of a geologic column from which a particular sequence of geologic processes may be inferred as described in the text. The section might be exposed, for example, as a roadcut, or as a cliff face on the side of a mountain. Adapted from Press and Siever (1978) by permission of W.H. Freeman and Company.



**Figure 8.1.** Age of Earth as estimated by various techniques since the Renaissance. Buffon tried to use rate of cooling of iron from a molten state to estimate the age of Earth, but the omission of as-yet undiscovered radioactive heating of Earth's interior (chapters 9 and 11) seriously shortened his cooling time. Joly worked out how long it would take to bring the oceans up to their current salinity based on the rate at which rivers carry salt to the sea; he ignored the precipitation of salt out of the ocean water into seafloor sediments. From Press and Siever (1978) by permission of W.H. Freeman and Company.



## Genesis

1:1 In the beginning God created the heaven and the earth.

1:2 And the earth was without form, and void; and darkness was upon the face of the deep. And the Spirit of God moved upon the face of the waters.

1:3 And God said, Let there be light: and there was light.

1:4 And God saw the light, that it was good: and God divided the light from the darkness.

1:5 And God called the light Day, and the darkness he called Night. And the evening and the morning were the first day.

- 11:10 These are the generations of Shem: Shem was an hundred years old, and begat Arphaxad two years after the flood:
- 11:11 And Shem lived after he begat Arphaxad five hundred years, and begat sons and daughters.
- 11:12 And Arphaxad lived five and thirty years, and begat Salah:
- 11:13 And Arphaxad lived after he begat Salah four hundred and three years, and begat sons and daughters.
- 11:14 And Salah lived thirty years, and begat Eber:
- 11:15 And Salah lived after he begat Eber four hundred and three years, and begat sons and daughters.
- 11:16 And Eber lived four and thirty years, and begat Peleg:
- 11:17 And Eber lived after he begat Peleg four hundred and thirty years, and begat sons and daughters.
- 11:18 And Peleg lived thirty years, and begat Reu:
- 11:19 And Peleg lived after he begat Reu two hundred and nine years, and begat sons and daughters.
- 11:20 And Reu lived two and thirty years, and begat Serug:
- 11:21 And Reu lived after he begat Serug two hundred and seven years, and begat sons and daughters.
- 11:22 And Serug lived thirty years, and begat Nahor:
- 11:23 And Serug lived after he begat Nahor two hundred years, and begat sons and daughters.
- 11:24 And Nahor lived nine and twenty years, and begat Terah:
- 11:25 And Nahor lived after he begat Terah an hundred and nineteen years, and begat sons and daughters.
- 11:26 And Terah lived seventy years, and begat Abram, Nahor, and Haran.



The year of the World.	2	<i>The first Age of the World.</i>	The Julian Period.	The year before Christ.
		<p>finished his work which he intended, he then rested from all labour, and blessing the seventh day, he ordained and consecrated the Sabbath, [Gen. 2. vers. 2, 3.] because therein he took breath, as himself is pleased to speak of himself, [Exod. 31. vers. 17.] and, as it were, refreshed himself. Nor as yet (for ought appeareth) had sin set footing into the world. Nor was there any punishment laid by God, either upon man-kinde, or upon Angels. Whence it was, that this afterward was set forth for a signe, as well of our sanctification in this world [Exod. 31. vers. 13.] as of that eternall Sabbath, to be enjoyed hereafter; wherein we expect a full deliverance and discharge from sin, and the dregs thereof, and all punishments belonging thereunto, [Heb. 4. vers. 4, 9, 10.]</p> <p>After the first week of the world ended, as it seemeth, God brought the new married couple into the garden of Eden, and charged them not to eat of the tree of knowledge of good and evil; but left them free use of all the rest.</p> <p>But the Devil, envying Gods honour and mans felicity, tempted the woman to sin by the Serpent; whence himself got the name and title of <i>the old Serpent</i>. [Apo. 12. vers. 9. and 20. vers. 2.] The woman then beguiled by the Serpent, and the man seduced by the woman, brake the ordinance of God concerning the forbidden fruit; and accordingly being called, and convicted of the crime, had their severall punishments inflicted on them: yet with this promise added, that the Seed of the woman should, one day, break the Serpents head, (i.) That Christ in the fulnesse of time should undo the works of the Devil, [1 Joh. 3. vers. 8. Rom. 16. vers. 20.] From whence it was, that Adam then first called his wife <i>Eve</i>; because she was then ordained to be the mother, not onely of all that should live this naturall life, but, of those also who should live by faith in her seed; which was the promised <i>Messias</i>: as <i>Sara</i> also afterward was counted the mother of the faithfull, [1 Pet. 3. verse 6. Gal. 4. verse 31.]</p> <p>Upon this occasion our first Parents, clad by God with raiment of skinnies, were turned out of <i>Eden</i>, and a fierie flaming sword set to keep the way leading to the tree of Life, to the end they should never alter eat of that fruit, which hitherto they had not touched [Gen. 3. verse 21, 22, &amp;c.] whence it is very probable, that Adam was turned out of Paradise the self same day that he was brought into it, which seemeth to have been upon the tenth day of the world (answering to our first day of <i>November</i>, according to supposition of the Julian Period) upon which day also, in remembrance of so remarkable a thing, as in all reason it should seem, was appointed the solemnity of <i>Expiation</i>, or attonement, and the yearly fast, spoken of by Saint Paul, <i>Acts 27. verse 9.</i> termed more especially by the name of <i>iasias</i>, wherein all, as well strangers as home born people, were commanded to afflict their souls with a most severe intermination, that every soul which should not afflict it self upon that day should be destroyed from among his people, [Lev. 16. v. 29. and 23. verse 29.]</p> <p>After the fall of Adam, <i>Cain</i> was the first of all mortall men that was born of a woman, [Gen. 4. verse 1.]</p> <p><i>Abel</i> being murdered by his brother <i>Cain</i>, the first born of all man-kind, God gave <i>Eve</i> another son in his stead; whence his name was called <i>Seth</i>, c. 4. v. 25. when Adam had now lived 130 years, c. 5. v. 3. From whence it is gathered, that between the death of <i>Abel</i>, and the birth of <i>Seth</i>, there was no other son born to <i>Eve</i>; for then he should have been recorded to have been given her instead of him: so that whereas now the race of man-kind had been continued to the terme of 128 years, it is probable, that the number of men was so encreased in the world, that <i>Cain</i> might justly fear, through the conscience of his crime, that every man that met him would also slay him. [c. 4. v. 14, 15.]</p> <p><i>Seth</i> now being 105 years old, begat a son, whom he named <i>Enoch</i>; which signifies, the lamentable condition of all mankind. For even then was the worship of God wretchedly corrupted by the race of <i>Cain</i>: whence it came, that men were even then so distinguished, that they who persisted in the true worship of God, were known by the name of the children of God; and they which forsook him, were termed the children of men, <i>Gen. 4. v. 26.</i> and 6. 1, 2.</p> <p><i>Cainan</i> the son of <i>Enoch</i> was born when his father was 90 years old, [c. 5. v. 10.]</p> <p><i>Mahaleel</i> was born when <i>Cainan</i> his father had lived 70 years, [c. 5. v. 12.]</p> <p><i>Jared</i> was born when his father <i>Mahaleel</i> had lived 65 years, [c. 5. v. 15.]</p> <p><i>Enoch</i> was borne when his father <i>Jared</i> had lived 162 years, [c. 5. v. 18.]</p> <p><i>Mathusalah</i> was born when <i>Enoch</i> his father had lived 65 years [c. 5. v. 25.]</p> <p><i>Lamech</i> was born when his father <i>Mathusalah</i> had lived 187 years, [c. 5. v. 25.]</p> <p>Now <i>Adam</i> the first father of all man kind, died when he had lived 930 years.</p> <p>As for <i>Enoch</i>, the seventh from <i>Adam</i>, God translated him in an instant, whiles he was walking</p>		
130.	d.			
235.	d.			945. 3769.
325.	d.			1035. 3679.
395.	d.			1015. 3069.
460.	d.			1017. 3544.
622.	d.			1332. 3383.
687.	d.			1327. 3317.
874.	d.			1584. 3130.
930.	d.			1640. 3074.
987.	d.			1697. 3017.

Fig. 2.3. Old Testament chronology published by Bishop James Ussher (1658).

In the Beginning, God created Heaven and Earth, Gen. I.V.I. Which beginning of time, according to our chronology, fell upon the entrance of the night preceding the twenty third day of October in the year of the Julian calendar, 710 (Ussher, 1658).

Thus, Ussher arrived at the beginning of night (evening), October 22, 4004 B.C., for the date of Creation, giving a universe 5,994 years old in 1990.



Basically, he counted begets since Adam & Eve.

4004 BC at the beginning of night (in Greenwich?) Oct 22!

~~All celebrated estimates of the age of the earth~~

Buffon - cooling of iron ~~ball~~ sphere

Comte de Buffon 1707 - 1788

one of most productive & best-known scientists of 18th century

laid the foundations for theory of probability

best known as a paleontologist

35 volume treatise Historie Naturelle

Had a whole theory - formed when a ~~comet~~ comet struck the sun - began as a white-hot globe - how long would it take to cool to present temperature?

Surface cool enough to touch

10 iron spheres

$\frac{1}{2}$ " 1" ~~1.5"~~  $1\frac{1}{2}$ " ... 5"

heated to white hot & observed  
time to cool to a touchable  
temperature

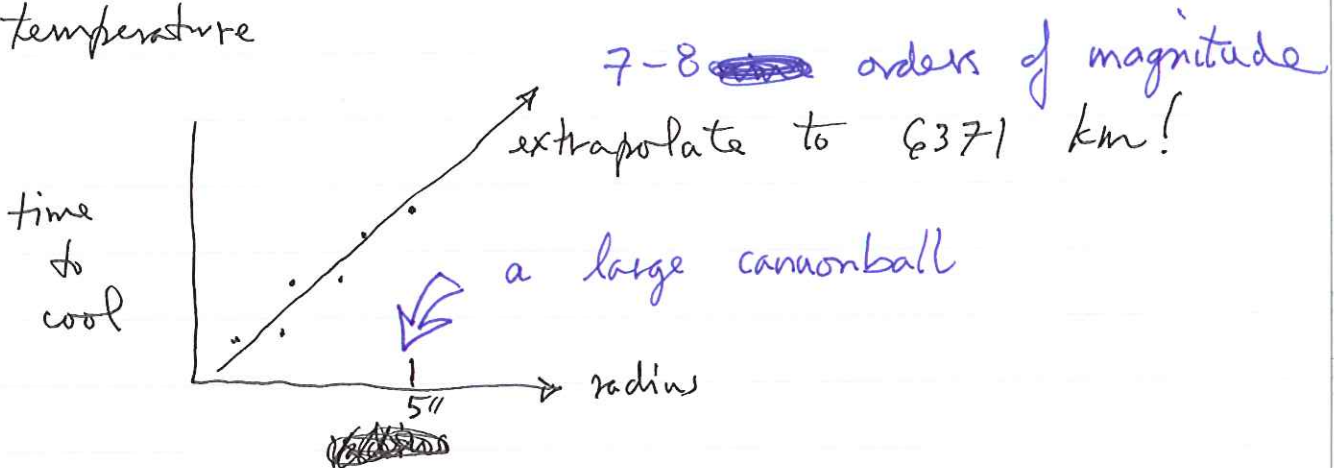


Table from Delrymple — note the  
air of exactitude 74,832 years

Clear from his writings, though, that  
he was skeptical of this result.

Was aware of geological evidence, e.g.  
thick layers of sediments exposed  
in Alps.

Lord Kelvin born 1824 Belfast

Greatest scientist of Victorian age  
Enunciated laws of thermodynamics  
Kelvin = degrees above absolute zero

Victorian  
equivalent of  
Einstein of  
the  
20th  
century



n globe. For more than half a century, thinking about the antiquity of the shafts, and to subsequently, are among attempts to determine a rational age for the Earth. But the concept that a date for the Earth could be determined from calculations of cooling was in evidence before.

Newton, English physicist and mathematician, was a leading figure of the Enlightenment. His theories of gravitation and inventions were the cornerstones of the scientific revolution. Newton was also interested in the cooling of heated bodies in air, he concluded that the rate of cooling was proportional to the difference in temperature between the body and the surrounding medium. Newton attempted to calculate an age for the Earth, but his calculations were flawed. In *Naturalis Principia Mathematica*, published in 1687, he speculated on the length of time it would take for the Earth to cool from iron the size of the Earth to cool:

...a globe, exposed red hot to the open air, will cool in a certain time; but a greater globe would retain its heat longer, because the surface (in proportion to the volume) is in that proportion as the diameter is to the cube of the diameter; and therefore a globe of iron, that is, about 40,000,000 feet in diameter, would require a number of days, or in above 50,000 years, to cool to the point of congelation; and I have investigated by experiments

...or at least suspected, that the relation between the time of cooling for bodies of planetary dimensions appeared at first glance. As we shall see, the relation was even more complicated than Newton's.

An analog for the cooling Earth was a material of everyday experience. Baron de Buffon (1707–1788), German philosopher and mathematician (and one of the early world, was an early subscriber to the *Philosophical Transactions of the Royal Society of London*. He proposed that the Earth

solidified in stages similar to those he had observed in the cooling of large masses of metal. The cooling Earth, according to Leibniz, was sculpted by large bubbles, some of which hardened into mountains while others collapsed to form valleys (Haber, 1959: 84–88). Like Newton, Leibniz did not venture to determine an age for the Earth from cooling. That bold step was left to another prominent figure of the Enlightenment.

Georges-Louis Leclerc, Comte de Buffon, was born September 7, 1707, at Montbard in Burgundy. He was educated in the law at the College of Godrans in Dijon and went on to study medicine, botany, and mathematics at Angers. Buffon was a man of enormous talent and energy. In addition to his scientific interests, he managed his family land holdings at Buffon and Montbard, engaged in the business of harvesting timber, established a commercial tree nursery, and built and operated an iron foundry, which played an important role in his research into the age of the Earth. His business activities, however, did not detract from his interest in science.

Buffon was one of the most productive and well-known scientists of the eighteenth century, and during his career he was elected to both the Royal Society and the Académie Française. He made fundamental contributions to the calculus of probability, plant physiology, and the scientific method, and laid the foundations for what would become the field of paleontology. He is best known, however, for an encyclopedic work in which he attempted, with considerable success, to synthesize all knowledge of nature and natural history into an intelligible whole. *Histoire Naturelle, Générale et Particulière* was originally intended to include an ambitious fifty volumes, of which Buffon actually completed 35 before his death in 1788. Among them were 12 volumes on mammals, nine on birds, and five on minerals as well as three introductory volumes and six lengthy supplements. The fifth supplement and the twentieth volume is *Epochs of Nature*, which was published in 1778 and is probably the best known. In it Buffon divided the history of the Earth into seven epochs. In the first epoch the Earth was a molten globe and the final epoch included the advent of man and the world as it is today; the intervening epochs included the formation of Earth's surface, the appearance of oceans and the beginnings of life, the formation of continents, the development of mammals, and the separation of the American and Eurasian continents.<sup>7</sup>

According to Buffon, the first epoch began when a comet collided with the Sun, causing the ejection of hot gases and liquid to

7. Both Albritton (1980: 84) and Haber (1959: 124–25) suggested that Buffon's use of seven epochs was a device to defuse church criticism of his concept of lengthy geological time, as the epochs are arbitrary and not an essential part of his synthesis of Earth history.

Buffon



form the planets of the Solar System and their satellites. Buffon's proposition that the Earth began as a molten globe was consistent with accepted cosmogonies of the time. Leibniz had argued for a liquid primitive Earth, and in *Principia* Newton had furnished proof that the Earth was once fluid and had cooled in the shape of an oblate spheroid, thus explaining the equatorial bulge. Buffon also leaned heavily on the work of Jean-Jacques Dortous de Mairan, a French pioneer in atmospheric physics known for his work on the aurora. De Mairan had compiled more than a half century of measurements on temperatures of the atmosphere, of hot springs, in mines, and on the formation of ice in surface waters. These measurements, published in 1749, convinced Mairan and Buffon, with whom Mairan corresponded, that the Earth contained residual heat and was cooling. The next logical step was to calculate the time required for the Earth to cool from its initial to its present state.

Rather than speculate, Buffon had his foundry fabricate ten iron spheres whose diameters varied in half-inch increments up to 5 inches. These he heated to white heat and then observed the time required for them to cool, first to red heat, then to absence of glow, then to a point where they were cool enough to hold in his hand, and finally to room temperature. To ensure uniform conditions and minimize the daily temperature fluctuations caused by the Sun, Buffon performed his experiments in a cellar laboratory. He found an approximately linear relation between diameter and cooling time, which he then logically but naively extrapolated to a sphere the size of the Earth. On this basis he calculated that a mass of molten iron the size of the Earth would require 42,964 years to cool below incandescence and 96,670 years to cool to the present temperature of the Earth.

Buffon then performed similar experiments on a second set of graduated spheres composed of materials nearer the actual composition of the Earth. He corrected his calculations for the delaying effect of the Sun's heat and combined these data with some major events in Earth's history as reconstructed in *Epochs* to deduce the following scale of times, each in years from the beginning (Haber, 1959: 118):

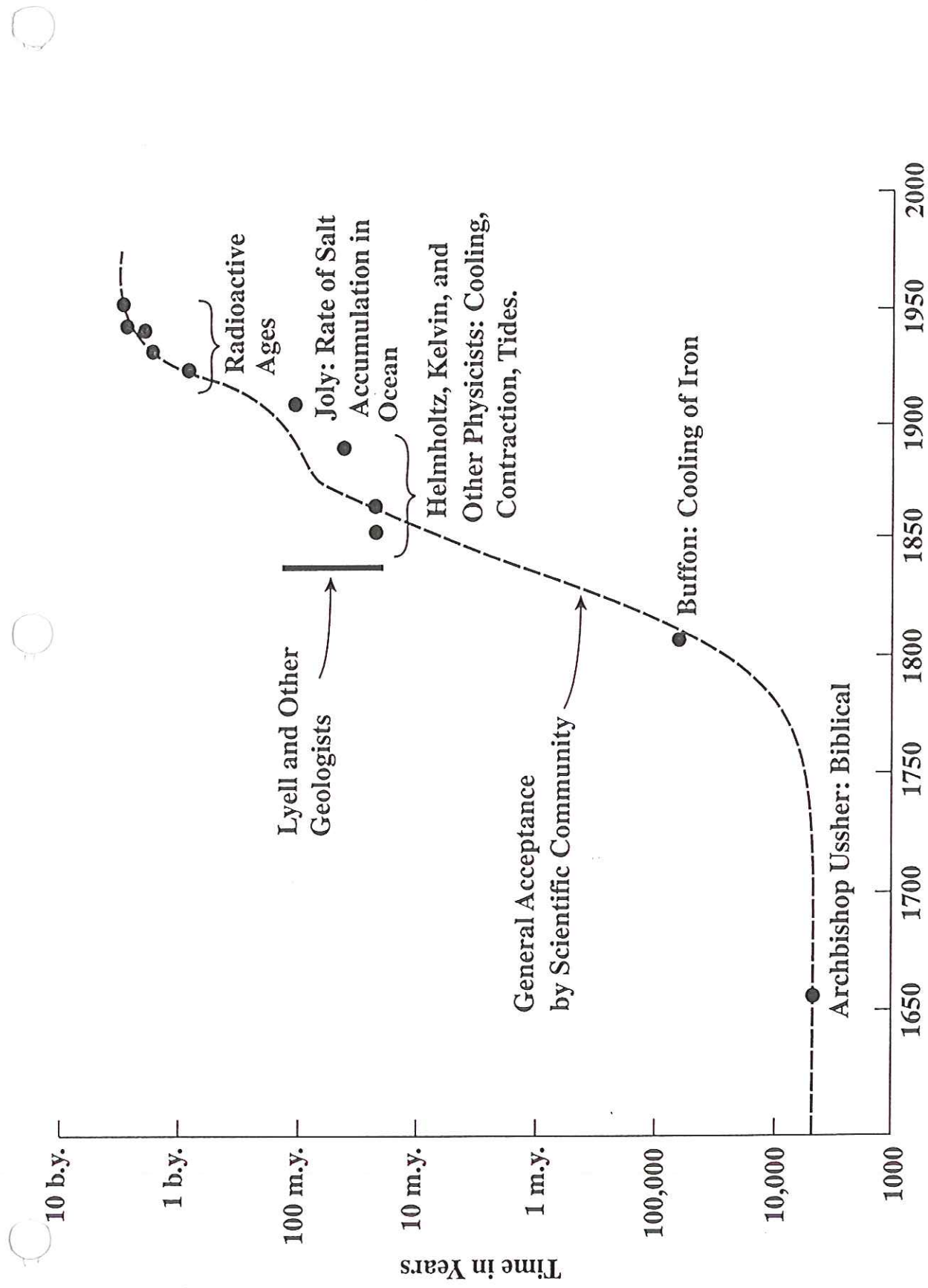
Surface of Earth consolidated	1
Earth consolidated to center	2,936
Earth cool enough to be touched	34,270
Beginning of life	35,983
Temperature of present reached	74,832
End of life	168,123

Although the calculations were detailed and carried an air of exactitude, Buffon was suspicious of his results because he felt that the



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Temperature of present reached	74,832
End of life	168,123 ← 700 co/ed for life



**Figure 8.1.** Age of Earth as estimated by various techniques since the Renaissance. Buffon tried to use rate of cooling of iron from a molten state to estimate the age of Earth, but the omission of as-yet undiscovered radioactive heating of Earth's interior (chapters 9 and 11) seriously shortened his cooling time. Joly worked out how long it would take to bring the oceans up to their current salinity based on the rate at which rivers carry salt to the sea; he ignored the precipitation of salt out of the ocean water into seafloor sediments. From Press and Siever (1978) by permission of W.H. Freeman and Company.



University of Glasgow at the age of ten, where he studied mathematics and became particularly interested in the work of Joseph Fourier, a French physicist who had devised mathematical methods of describing the physics of heat. In 1841 Thomson entered Cambridge University, from which he graduated in 1845 with high honors. Following graduation, he spent several months at the University of Paris, where he gained experience in laboratory methods by working with Henry Regnault, the French physicist. In 1846, largely through the efforts of his father, Thomson was appointed Professor of Natural Philosophy at Glasgow, a chair he held throughout his long and distinguished career.

Thomson was an unusually productive scientist, and at the time of his retirement in 1899 he had authored more than 600 scientific papers and books on electricity, magnetism, *thermodynamics*, hydrodynamics, atmospheric electricity, geomagnetism, *geodesy*, the thermal state and rotation of the Earth, tidal theory, and the age of the Earth. In 1848 he devised and described the absolute temperature scale, which still carries his name (the *Kelvin scale*). One of his most significant scientific achievements, a statement of the second law of thermodynamics, was published in 1851 under the title "On the Dynamical Theory of Heat." The cosmic implications of this fundamental principle later formed the basis for his calculations of the age of the Earth.

Kelvin

Thomson's work on electricity led him into the field of applied physics and to a role in the laying of the first trans-Atlantic telegraph cable. He became a director of the Atlantic Telegraph Company in 1856, served as the electrician aboard the *Agamemnon* in the first unsuccessful attempt to lay an Atlantic cable in 1858, and supervised the successful laying of the first cable by the *Great Eastern* in 1866, for which he was knighted that year by Queen Victoria. Thomson was also a prolific inventor, being responsible for the mirror galvanometer and siphon recorder used to receive telegraph signals, the stranded electrical conductor, the tide gauge, and an improved mariner's compass that allowed compensation for the magnetism of a steel ship. By the end of his career he held some 70 patents.

Thomson was probably the most honored British scientist in history, with countless awards, medals, and degrees. He was first elected president of the Royal Society in 1890 and was reelected to that office for five consecutive terms. In 1882 he was raised to the peerage and became Baron Kelvin of Largs, Ayrshire, and in 1896, on the fiftieth anniversary of his professorship at Glasgow, he was awarded the Grand Cross of the Royal Victorian Order. Because of his renown and the great respect in which he was held throughout the Western world, Kelvin's writings and opinions were accorded special significance. It



Also eminently practical — lead from  
Dalrymple

- first submarine telegraph cable
- improved mariners' compass
- tide gauge

Estimated age of Sun — knew the  
rate at which it radiates energy in  
form of light waves

$$4 \cdot 10^{26} \text{ W} \text{ — ignores neutrinos}$$

At the time only two known sources  
of energy available to power the sun

- chemical combustion, as in  
a coal furnace

Combustion of one kg of coal  
releases  $3 \cdot 10^7 \text{ J}$

$$\text{coal: } 3 \cdot 10^7 \text{ J/kg}$$

For how long could the sun burn  
if its energy were supplied by  
chemical processes, i.e. suppose the  
Sun is a coal furnace



similar to homework problem

$$\underbrace{4 \cdot 10^{26} \frac{\text{J}}{\text{sec}}}_{\text{heat release rate}} \times t \text{ (sec)} = \underbrace{2 \cdot 10^{30} \text{ kg}}_{\text{mass of Sun}} \times \underbrace{3 \cdot 10^7 \frac{\text{J}}{\text{kg}}}_{\text{released by burning}}$$

$$t = \frac{(2 \cdot 10^{30}) (3 \cdot 10^7)}{4 \cdot 10^{26}} = 1.5 \cdot 10^{11} \text{ sec}$$

one year  $\approx 3 \cdot 10^7$  sec

$$t = 4800 \text{ years}$$

Pretty close to Archbishop Ussher!

~~Other ideas better than this, can be proposed that the source of the Sun's heat was the only other than known process.~~

~~meteoric aggregations and feeding of mass to the Sun by the influx of meteorites & comets~~

- meteoric aggregation — i.e. constant infall of meteorites into Sun

Kelvin also eliminated this on the basis of detailed observations of the orbit of Mercury

He thus concluded that the Sun had to have formed hot, and that it is cooling off — all the radiated energy is simply this fossil initial heat

To estimate age of Sun from cooling need to know two things:

- ~~average~~ average temperature of Sun  $T = 15,000^\circ\text{C}$
- specific heat of solar material

$$c = \frac{J}{\text{kg}^\circ\text{C}}$$

amount of heat needed to raise  $T$  by  $1^\circ\text{C}$  — or amount that must be removed to cool it by  $1^\circ\text{C}$



↖ actually 4184

$$c(H_2O) \cong \frac{4000 \text{ J}}{\text{kg } ^\circ\text{C}}$$

If sun ~~was~~ has same specific heat of water, how much is cooling off?

$$\frac{c_{\text{total } \odot} \quad \swarrow \text{ in one year}}{c \left( \frac{\text{J}}{\text{kg}} \right) M_{\odot} (\text{kg}) \quad dT} = 4 \cdot 10^{26} \frac{\text{J}}{\text{sec}}$$

$$dt (3 \cdot 10^7 \text{ sec} = 1 \text{ yr})$$

or, better

$$\underbrace{\left( 4 \cdot 10^{26} \frac{\text{J}}{\text{sec}} \right) \left( 3 \cdot 10^7 \frac{\text{sec}}{\text{year}} \right)}_{\text{heat lost in 1 yr}} = c \left( \frac{\text{J}}{\text{kg}} \right) M_{\odot} (\text{kg}) dT$$

$$dT = \frac{(4 \cdot 10^{26})(3 \cdot 10^7)}{(4000)(2 \cdot 10^{30})}$$

if  $c_{\odot} = c_{H_2O} \rightarrow$   $dT = 1.5^\circ\text{C / year}$

This would result in a thermal contraction

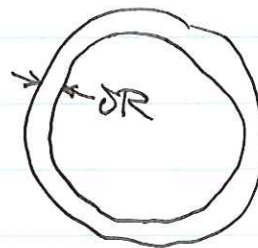
~~$\alpha_{H_2O} = 3.6 \cdot 10^{-4} \text{ (C)}^{-1}$~~

~~$\frac{\Delta R}{R} = 3.6 \cdot 10^{-4} \text{ per } ^\circ\text{C cooling}$~~

~~$= 1.5 \cdot 10^{-4} \text{ per year}$~~

$\alpha_{H_2O} \approx 10^{-4} \text{ per } ^\circ\text{C}$

$\frac{\Delta R}{R} = 10^{-4} \text{ per } ^\circ\text{C}$   
 $= 1.5 \cdot 10^{-4} \text{ per year}$



$= 10\% \text{ shrinking in 70 years}$

Remember:  
 Kelvin knew  
 the mean  
 density of  
 the Sun  
 from  
 Kepler's 3rd law

$\rho_{\odot} = 1400 \text{ kg/m}^3$

This could clearly be ruled out by observations.

Concluded that

$c_{\text{solar material}} = 10^2 - 10^4 \times c_{H_2O}$

Say  $c_{\text{sun}} = 1000 c_{H_2O} = 4 \cdot 10^6 \frac{\text{J}}{\text{kg}}$

Then cooling rate  $1.5 ^\circ\text{C} / 1000 \text{ years}$

Age  $t$ :  $\frac{c M_{\odot} 15,000 ^\circ\text{C}}{1.5 ^\circ\text{C} / 10^4 \text{ yr}} = 4 \cdot 10^{28} \frac{\text{J}}{\text{sec}}$

$t \text{ (sec)}$

varies do to that  
 rate to that  
 $10^4 \text{ yr} / 10^4 \text{ yr}$   
 $10^4 \text{ yr} / 10^4 \text{ yr}$



$$t = \frac{\left(4 \cdot 10^6 \frac{\text{J}}{\text{kg} \cdot ^\circ\text{C}}\right) \left(\cancel{100} 2 \cdot 10^{30} \text{ kg}\right) \left(15,000 ^\circ\text{C}\right)}{4 \cdot 10^{26} \text{ J/sec}}$$

$$= 3 \cdot 10^{14} \text{ sec} = \boxed{110 \text{ m.y.}}$$

Read  $\Pi$  from Dalrymple  $\leftarrow$  go to page 9 1/2

We now know, of course, that the Sun's energy is provided by nuclear burning — see homework problem

John Joly 1857 — 1933 Geology  
 professor at U. Dublin — best  
 known for his work on the  
 chemical extraction of radium &  
 its use in treatment of cancer

$$\text{age of } \oplus = \frac{\text{total Na}^+ \text{ in ocean}}{\text{annual Na}^+ \text{ influx from rivers}}$$

It seems, therefore, on the whole most probable that the sun has not illuminated the earth for 100,000,000 years, and almost certain that he has not done so for 500,000,000 years. As for the future, we may say, with equal certainty, that inhabitants of the earth cannot continue to enjoy the light and heat essential to their life, for many million years longer, unless sources now unknown to us are prepared in the great storehouse of creation (W. Thomson, 1862a: 494).

This last statement was prophetic. There were indeed powerful and unknown sources of energy fueling the Sun's fires, but their discovery was still four decades in the future.



A few years later Kelvin returned to the problem.

Used observations of the rate of temp increase in mines — cooling of  $\oplus$  rather than sun

Do not discuss here since need to introduce concepts of heat flow and thermal conductivity

On this basis he obtained 98 m.y. In good agreement with his solar estimate but he felt it was considerably more certain — this in fact his most famous estimate.

His 98 Myr age assumed a very high ~~temperature~~ (molten)  $T_0 = 3870^\circ\text{C}$

Clarence King (Director of USGS) improved upon this — ~~the~~ evidence for a much lower  $T_0 \approx 1200^\circ\text{C}$

This lowered the age to 20-40 Myr



seems unlikely that the Moon was ever closer to the Earth than about 38 Earth radii (compared to the present 60).

Within the framework of late-nineteenth-century understanding of the Earth, Darwin's hypothesis was reasonable and appropriate, but we now know that tidal considerations are incapable of providing a basis for determining the Earth's age. One has the feeling that Sir George wouldn't have minded.

### The Salt Clock

Physics was not the only discipline to provide early estimates of Earth's age; chemistry was also set to the task, and in a rather clever way.

Imagine a tub of water to which a chemical is continuously added. If you knew the amounts of the chemical in the water both now and when the tub was first filled, and the rate at which the chemical was added to the water, then you could calculate the time of origin, i.e. the age, of the tub of water. It was the possibility of just such a calculation that Edmund Halley (1656–1742), the Astronomer Royal who predicted the return of the comet that bears his name, had in mind when, in 1715, he proposed that the age of the Earth might be calculated from the salt content of the ocean and of certain kinds of lakes (Halley, 1715).<sup>19</sup>

Halley observed that all lakes that receive runoff from rivers but lack outflow contain salt in varying amounts. The concentration of salt in the waters of these lakes must increase, he said, because salt, picked up by the rivers in their passage over the Earth, is continuously added but not removed.<sup>20</sup> Water is removed by evaporation, but water vapor is fresh so the salt is left behind: "But the vapours thus exhaled are perfectly fresh; so that the saline particles brought in by the rivers remain behind, while the fresh evaporates; and hence it is evident that the salt in the lakes will be continually augmented, and the water grow salter and salter. . . ." (Halley, 1715, in G. F. Becker, 1910a: 460). If this was truly the cause of the saltiness of lakes, Halley reasoned, then it is probable that the same mechanism was responsible for the saltiness of the ocean.

Analytical methods were then incapable of measuring the minute quantities of salt in the rivers that supplied the ocean, but Halley

19. Lengthy excerpts from Halley's paper appear in G. F. Becker (1910a).

20. The idea that the ocean was originally fresh and that its salt was dissolved out of the Earth's crust did not originate with Halley. Leibniz, among others, had incorporated the idea into his model of the Earth.



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had another method in mind. If the concentration of salt in the ocean  
was measured at different times, then the rate of addition and the age  
of the ocean, which he equated with the age of the Earth, could be  
determined. A repeat of the experiment at some later date, he noted,  
would not only check the constancy of the rate at which salt is added  
but would verify the hypothesis.

Halley observed, however, that the age of the Earth could not  
yet be calculated. The experiment would require a very long time, and  
he lamented that the ancient Greek and Latin authors had not pro-  
vided information on the saltiness of the ocean 2,000 years ago. The  
only thing that could be done, advised Halley, was for the Royal So-  
ciety to ensure that future generations would be provided with the  
necessary data: "I recommend it therefore to the society, as opportu-  
nity shall offer, to procure the experiments to be made of the present  
degree of saltiness of the Ocean, and as many of these lakes as can be  
come at, that they may stand upon record for the benefit of future  
ages" (Halley, 1715, in G. F. Becker, 1910a: 461). Halley also observed  
that the method would provide only a maximum age for the Earth,  
because the ocean and some lakes might well have contained some  
salt when they first formed. The experiment, he said, was still worth-  
while because "[it] is chiefly intended to refute the ancient notion,  
some have of late entertained, of the eternity of all things; though per-  
haps by it the world may be found much older than many have hith-  
erto imagined" (Halley, 1715, in G. F. Becker, 1910a: 461).

Halley's idea was never pursued and seems to have been largely  
forgotten until 1876, when T. Mellard Reade rediscovered the method  
he called "chemical denudation." Reade proposed that the age of the  
ocean could be found from the concentrations of *chlorides* and *sulfates*.  
Instead of determining the rates of addition of these compounds by  
measuring their concentrations at different times, Reade proposed to  
find the values by estimating the annual amounts carried into the  
ocean by the rivers of the world. At the present annual rates of addi-  
tion, he calculated that it would require 25 Ma for the sulfates of cal-  
cium and magnesium to reach their present concentrations in ocean  
water; for chlorides (principally of sodium) the comparable time was  
200 Ma (Reade, 1876, 1879).

Reade's basic idea of dating the Earth from the progressive  
change in the chemistry of the ocean was carried to a high degree of  
refinement by John Joly (1857–1933), a professor of geology and min-  
eralogy at the University of Dublin. The son of a clergyman, Joly was  
born in Holywood, Kings County, Ireland, and educated at Trinity  
College, Dublin, where he taught throughout his career. Joly was



trained as a physicist but developed a keen interest in geology, an evolution that is reflected in his sequence of positions at Trinity College, where he held appointments as demonstrator in civil engineering (1883), in physics (1893), and finally as professor of geology and mineralogy (1897).

Joly's accomplishments were numerous. He developed a method of extracting radium and pioneered its use in the treatment of cancer. He invented a type of thermometer, a steam calorimeter to measure heat energy, and a photometer to measure light frequencies. He was the first to propose that convection, driven by the heat from radioactive decay in the Earth's interior, might play a major role in the energetics and evolution of the Earth's crust. During his distinguished career he was accorded many honors, including election as a Fellow of the Royal Society in 1892.

Joly's (1899) classic paper "An Estimate of the Geological Age of the Earth" was read to the Royal Dublin Society, of which Joly was then Secretary, on May 17, 1899. In it (p. 249) Joly proposed to measure the age of the Earth from the accumulation, not of a salt, but of an element in the waters of the ocean:

Now, if any of the elements entering the ocean is not again withdrawn, but is, in a word, "trapped" therein, reappears as no extensive marine deposit, and is not laid down sensibly upon its floor, and if the amount of uniformity already defined is accepted, evidently in the rate of annual accretion by the ocean, from the rivers, of this substance and the amount of it now in the ocean, the whole period since the beginning of its supply can be estimated.

Such an element, he said, was sodium, and by using the pure element he avoided the questions and uncertainties of ionization and chemical form.

The result of Joly's calculations was an age for the Earth that differed little from Kelvin's:

The quantity of sodium now in the sea, and the annual rate of its supply by the rivers, lead, it will be seen, to the deduction that the age of the earth is  $99 \times 10^6$  years. Certain deductions from this are, it will be shown, warranted, so that the final result of this paper will be to show that the probable age is about  $89 \times 10^6$  years. Also, that this is probably a major limit, and that considerable departure from uniformity of activities could hardly amend it to less than  $80 \times 10^6$  years (Joly, 1899: 249).

How did Joly arrive at these values? The basic equation, lacking certain necessary corrections, is the soul of simplicity:

$$\text{age of Earth} = \frac{\text{total sodium in ocean}}{\text{annual sodium influx from rivers}}$$



developed a keen interest in geology, an  
sequence of positions at Trinity Col-  
lege as demonstrator in civil engineer-  
ing and finally as professor of geology and

was numerous. He developed a method  
and used it in the treatment of cancer.  
In 1867, a steam calorimeter to measure  
heat, and a method to measure light frequencies. He was  
invention, driven by the heat from radio-  
activity, might play a major role in the en-  
richment of the earth's crust. During his distinguished  
career, including election as a Fellow

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the values? The basic equation, lacking  
the soul of simplicity:

total sodium in ocean  
sodium influx from rivers

Joly began by arguing that the extrapolation of the rates of present  
processes into the past was warranted unless it could be shown that  
the rates had been interrupted by catastrophe or change. He pointed  
out that the approximate constancy of erosion of the land surface  
throughout geological time was a tenet whose validity had not seri-  
ously been questioned. Even so, his calculation required acceptance  
of this tenet only in part:

... that part of it which refers to the removal of the land surface by solution.  
It has to be accepted as a preliminary step that this, on the whole, has been  
constant. Herein are involved a constancy, within certain fairly wide limits, of  
rainfall over the land areas; a constancy, within fairly wide limits (which can  
be roughly defined), of the exposed land area, and a constancy in the nature  
and rate of solvent actions going on over the land surfaces (Joly, 1899: 248).

The other tenet that must be accepted was that the primeval ocean did  
not contain sodium in the quantities now observed.

The determination of a value for the sodium in the modern  
ocean was not difficult. Sir John Murray, a British oceanographer and  
founder of the study of submarine geology, had made estimates of the  
mass and mean depth of the ocean, the total volume of river discharge,  
and the quantity of dissolved matter in a number of the world's rivers.<sup>21</sup>  
Joly used Murray's mean ocean depth of 3,797 m combined with Her-  
mann Wagner's ocean area of  $1.0372 \times 10^8 \text{ km}^2$  and the density of sea  
water to calculate that the mass of water in the world's ocean was  
 $1.3245 \times 10^8 \text{ metric tons}$ .<sup>22</sup> The salinity of the ocean was known to be  
3.5%, of which sodium chloride, NaCl, constitutes 77.758% of the  
total salts. Since 39.32% by weight of NaCl is Na, Joly arrived at a  
value of  $1.4177 \times 10^{16} \text{ metric tons}$  for the total mass of Na in the ocean.

Murray's estimate for the total annual discharge of rivers into the  
ocean was  $2.7176 \times 10^4 \text{ km}^3$ . His estimate for the dissolved matter in  
river water was based on analyses of waters from 19 rivers of the  
world. These analyses showed that of the total salts in river water,  
three contain sodium, giving a combined mass of 5,250 metric tons of  
Na per cubic kilometer of river water. The annual discharge multi-  
plied by the sodium content provided Joly with a value of  $1.4268 \times 10^8$   
metric tons of Na supplied per year by rivers to the ocean. Joly's un-  
corrected age for the Earth thus was

$$\text{age of Earth} = \frac{1.4177 \times 10^{16} \text{ metric tons Na}}{1.4268 \times 10^8 \text{ metric tons Na/yr}} = 99.4 \times 10^6 \text{ yr}$$

21. John Murray (1841–1914) was one of the organizers of the famed *Challenger*  
expedition of 1872–76 and served on board as a naturalist. He completed the report of  
the expedition after the death of the leader, Sir Wyville Thomson.

22. Joly's 1899 data were expressed in English units of measurement. I have con-  
verted them to metric units.



But this was not the true age of the Earth; some corrections were necessary.

Joly applied a correction to the numerator for the amount of sodium in the original ocean. This correction was based on his estimate of the amount of sodium that would have been dissolved out of the primeval rocks as the crust and ocean were forming. Like Kelvin, Joly presumed that the Earth began as a molten globe. At a temperature of  $1,500^{\circ}\text{C}$ , the crust would have been molten and the material of the future ocean would have consisted primarily of free hydrogen, free oxygen, and HCl gas; NaCl would not exist at that temperature. As the crust and primitive atmosphere cooled, first water vapor then liquid water would form and the hot acidic rains (from the HCl dissolved in the rain) would react with the hot crustal rocks to form salts of Na, K, Mg, Ca, and Fe. From the abundances of these elements in what Joly took to be the average crustal rock, and presuming that all of the HCl would be neutralized by reactions with these elements in the crust, he concluded that only 14% of the Cl in the acidic rains, i.e. in the original ocean, would combine with Na to form NaCl.

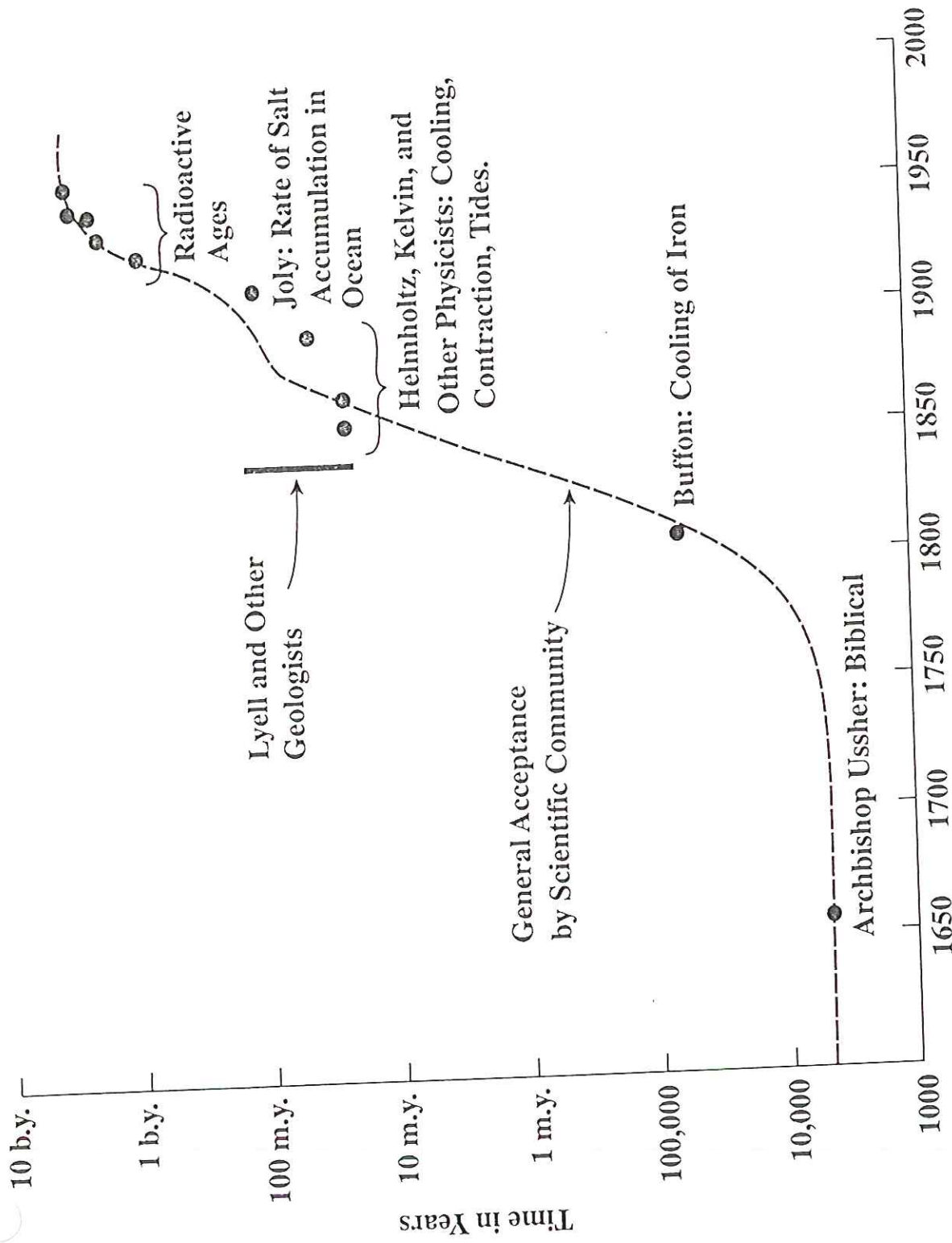
The next quantity Joly had to find was the total amount of Cl available in the primeval ocean. To do this Joly assumed that the amount of Cl now in the ocean is the sum of the amount in the primeval ocean and the amount brought in by rivers since the ocean first formed. The amount of Cl in the present ocean was relatively easy to estimate; it was the sum of the amounts in the ocean's NaCl ( $21.913 \times 10^{15}$  tons) and  $\text{MgCl}_2$  ( $3.775 \times 10^{15}$  tons).

The amount brought in by rivers over geologic time was a less certain calculation. Joly summed the amounts brought in each year as NaCl, LiCl, and  $\text{NH}_4\text{Cl}$ , but reduced the contribution from NaCl by 10% to account for recycling from the sea by evaporation. This sum, less the 10% correction, was about  $69 \times 10^6$  tons of Cl per year. But for how many years? Joly chose 86 Ma as a reasonable value, which yielded  $5.929 \times 10^{15}$  tons of Cl added to the ocean by rivers since the world began. To summarize, Joly's estimate of the Cl in the primeval ocean was

now as NaCl	$21.913 \times 10^{15}$ tons
now as $\text{MgCl}_2$	$3.775 \times 10^{15}$ tons
from rivers in $86 \times 10^6$ years	$-5.929 \times 10^{15}$ tons
Cl in primeval ocean	$19.759 \times 10^{15}$ tons

Fourteen percent of this Cl, or  $2.766 \times 10^{15}$  tons, would combine with  $1.789 \times 10^{15}$  tons of Na, which was Joly's value for the original Na in the primeval ocean and was the amount to be subtracted from the numerator in his age equation.

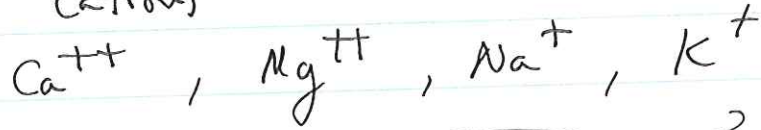




**Figure 8.1.** Age of Earth as estimated by various techniques since the Renaissance. Buffon tried to use rate of cooling of iron from a molten state to estimate the age of Earth, but the omission of as-yet undiscovered radioactive heating of Earth's interior (chapters 9 and 11) seriously shortened his cooling time. Joly worked out how long it would take to bring the oceans up to their current salinity based on the rate at which rivers carry salt to the sea; he ignored the precipitation of salt out of the ocean water into seafloor sediments. From Press and Siever (1978) by permission of [unclear] Company.

Table 8.1 dissolved species in sea water — except in estuaries very constant

Main cations



derived from weathering of feldspars and other minerals

Total dissolved solids 35 g/kg

Salinity = 35 ‰ (per mille)

How much  $\text{Na}^+$  in oceans?

$$\begin{aligned} \text{Volume} &= 4\pi (6371)^2 (0.7) \text{ km}^3 \quad \leftarrow \text{mean depth} \\ &= 1.4 \cdot 10^9 \text{ km}^3 \\ &= 1.4 \cdot 10^{27} \text{ liters} \end{aligned}$$

$$\begin{aligned} \# \text{ moles } \text{Na}^+ &= (1.4 \cdot 10^{27} \text{ l}) \left( 0.47 \frac{\text{moles}}{\text{l}} \right) \\ &= 6.7 \cdot 10^{20} \text{ moles } \text{Na}^+ \text{ in ocean} \end{aligned}$$

To find riverine influx need to measure average  $\text{Na}^+$  conc. in river water x total discharge



Fig. 5.3 & Table ~~5.3~~ <sup>5.4</sup> show average composition of dissolved species in river water

cations:  $\text{Ca}^{++}$ ,  $\text{Mg}^{++}$ ,  $\text{Na}^+$   
 anions:  $\text{Cl}^-$ ,  $\text{SO}_4^-$ ,  $\text{HCO}_3^-$

Note the effect of road salt in winter

NA natural	5.5 mg/l	} in NA rivers
polluted	8.4 mg/l	

Also  $\text{Cl}^-$  same reason

$\text{SO}_4^-$  pollution due to fertilization of fields — sulfuric acid used in fertilizer production  
 Also coal combustion

$\text{Na}^+$  in average world rivers

5.2 mg/l

$$\frac{(5.2 \cdot 10^{-3} \text{ g/l})}{23 \text{ g/mole}} = 2.3 \cdot 10^{-4} \text{ mole/l}$$

2000 times less than seawater

Table 5.1 lists major rivers in order of discharge  $\text{km}^3/\text{year}$   $\text{H}_2\text{O}$

Total all rivers  $37,400 \text{ km}^3/\text{yr}$

Amazon alone  $1/5$  of this

$$37,400 \text{ km}^3/\text{yr} = 3.7 \cdot 10^{16} \text{ l/yr}$$

$\text{Na}^+$  addition:

$$\left( 3.7 \cdot 10^{16} \text{ l/yr} \right) \left( 2.3 \cdot 10^{-4} \text{ moles/l} \right)$$

$$= 8.5 \cdot 10^{12} \text{ moles Na}^+ \text{ added/yr}$$

$$\text{Age of } \oplus = \frac{6.7 \cdot 10^{20} \text{ moles}}{8.5 \cdot 10^{12} \text{ moles/yr}}$$

$$= 79 \text{ million years}$$

Joly used somewhat different data and found  $99 \times 10^6$  years

"Certain deductions..." — he applied a small correction for initial  $\text{Na}^+$  in oceans

$$t_{\text{Joly}} = 89 \times 10^6 \text{ years}$$



**TABLE 8.1** Major Dissolved Components of Seawater for a Salinity of 35‰

Ion	Concentration		Percent Free Ion
	g/kg	mM <sup>a</sup>	
Cl <sup>-</sup>	19.354	558	100
Na <sup>+</sup>	10.77	479	98
Mg <sup>++</sup>	1.290	54.3	89
SO <sub>4</sub> <sup>--</sup>	2.712	28.9	39
Ca <sup>++</sup>	0.412	10.5	99
K <sup>+</sup>	0.399	10.4	98
HCO <sub>3</sub> <sup>-b</sup>	0.12	2.0	80

<sup>a</sup> mM = millimoles per liter at 25°C.

<sup>b</sup> For pH = 8.1, P = 1 atm, T = 25°C.

Sources: Wilson 1975; Skirrow 1975; Millero and Schreiber 1982.

**TABLE 5.6** Chemical Composition of Average River Water

By Continent	River Water Concentration <sup>a</sup> (mg/l)							Water			
	Ca <sup>++</sup>	Mg <sup>++</sup>	Na <sup>+</sup>	K <sup>+</sup>	Cl <sup>-</sup>	SO <sub>4</sub> <sup>--</sup>	HCO <sub>3</sub> <sup>-</sup>	SiO <sub>2</sub>	TDS	Discharge (10 <sup>3</sup> km <sup>3</sup> /yr)	Runoff Ratio <sup>b</sup>
Africa:											
Actual	5.7	2.2	4.4	1.4	4.1	4.2	26.9	12.0	60.5	3.41	0.28
Natural	5.3	2.2	3.8	1.4	3.4	3.2	26.7	12.0	57.8		
Asia:											
Actual	17.8	4.6	8.7	1.7	10.0	13.3	67.1	11.0	134.6	12.47	0.54
Natural	16.6	4.3	6.6	1.6	7.6	9.7	66.2	11.0	123.5		
S. America:											
Actual	6.3	1.4	3.3	1.0	4.1	3.8	24.4	10.3	54.6	11.04	0.41
Natural	6.3	1.4	3.3	1.0	4.1	3.5	24.4	10.3	54.3		
N. America:											
Actual	21.2	4.9	8.4	1.5	9.2	18.0	72.3	7.2	142.6	5.53	0.38
Natural	20.1	4.9	6.5	1.5	7.0	14.9	71.4	7.2	133.5		
Europe:											
Actual	31.7	6.7	16.5	1.8	20.0	35.5	86.0	6.8	212.8	2.56	0.42
Natural	24.2	5.2	3.2	1.1	4.7	15.1	80.1	6.8	140.3		
Oceania:											
Actual	15.2	3.8	7.6	1.1	6.8	7.7	65.6	16.3	125.3	2.40	—
Natural	15.0	3.8	7.0	1.1	5.9	6.5	65.1	16.3	120.3		
World average:											
Actual	14.7	3.7	7.2	1.4	8.3	11.5	53.0	10.4	110.1	37.4	0.46
Natural (unpolluted)	13.4	3.4	5.2	1.3	5.8	8.3	52.0	10.4	99.6	37.4	0.46
						(5.3) <sup>c</sup>	1.0	0	10.5	—	—
Pollution	1.3	0.3	2.0	0.1	2.5	3.2	2%	0%	—	—	—
						(6.2) <sup>c</sup>	28%	2%	—	—	—
World % pollutive	9%	8%	28%	7%	30%	28%	(54%) <sup>c</sup>	0%	—	—	—

<sup>a</sup> Actual concentrations include pollution. Natural concentrations are corrected for pollution.

<sup>b</sup> Runoff ratio = average runoff per unit area/average rainfall (calculated from Meybeck).

<sup>c</sup> We have raised pollutive contribution; see Table 5.11. (Our values are in parentheses.)

Source: All river water concentrations and discharge values from Meybeck (1979) except "actual" concentrations by continent, which were calculated from Meybeck's data. (M. Meybeck, "Concentrations des eaux fluviales en éléments majeurs et apports en solution aux océans," *Rev. Géol. Dyn. Geogr. Phys.*, 21(3), 220, 227.

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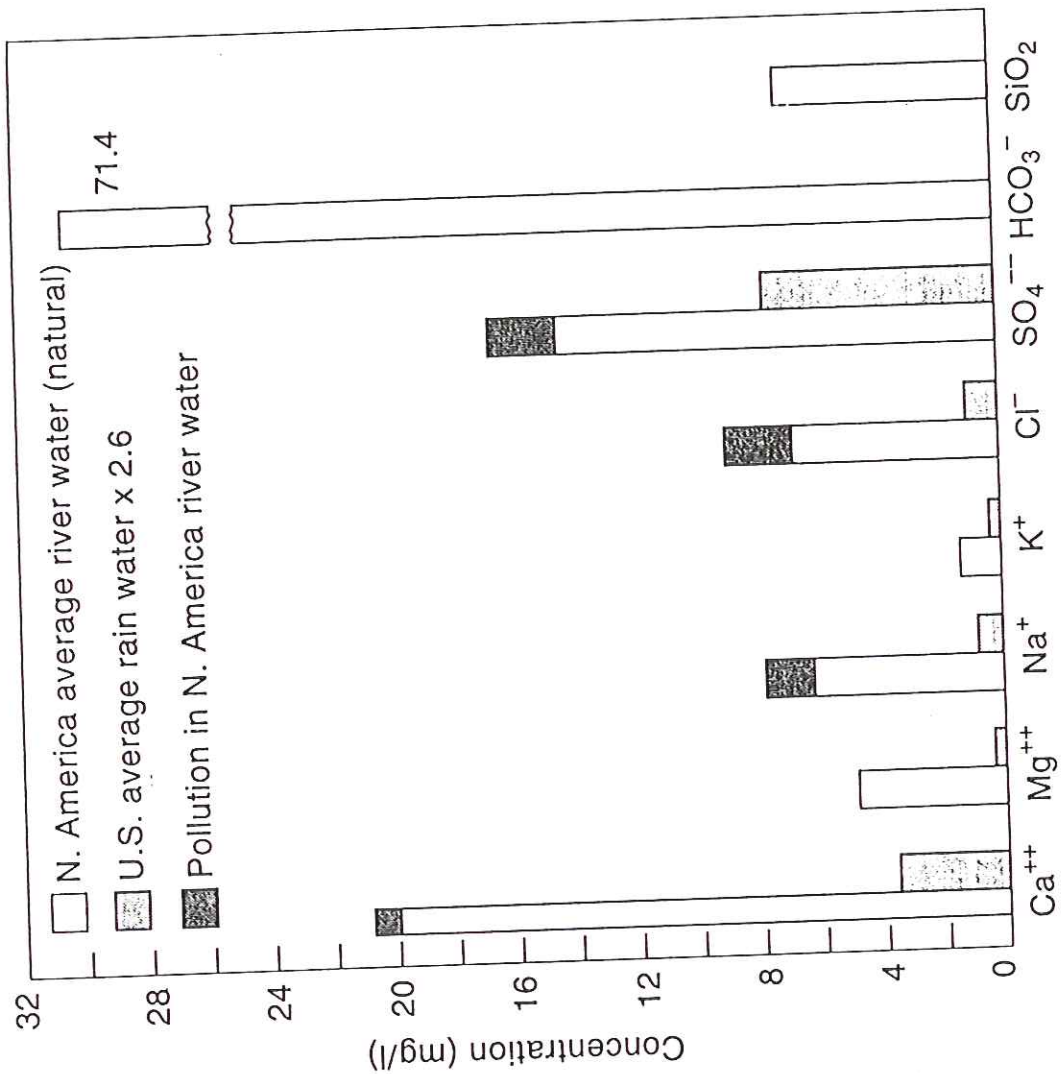


Figure 5.3. Comparison of dissolved composition of North American natural and polluted river water (data from Meybeck 1979) with U.S. rainwater (concentrations in mg/l). Rainwater concentrations are multiplied by 2.6 to correct for evaporation from the continents (see text).

**TABLE 5.1** Major Rivers that Flow to the Sea, Listed in Order of Discharge

River	Location	Water (km <sup>3</sup> /yr)
1. Amazon	S. America	6300
2. Zaire (Congo)	Africa	1250
3. Orinoco	S. America	1100
4. Yangtze (Chiang)	Asia (China)	900
5. Brahmaputra	Asia	603
6. Mississippi	N. America	580
7. Yenisei	Asia (Russia)	560
8. Lena	Asia (Russia)	525
9. Mekong	Asia (Vietnam)	470
10. Ganges	Asia	450
11. St. Lawrence	N. America	447
12. Parana	S. America	429
13. Irrawaddy	Asia (Burma)	428
15. Mackenzie	N. America	306
17. Columbia	N. America	251
20. Indus	Asia (India)	238
Red (Hungho)	Asia (Vietnam)	123
Huanghe (Yellow)	Asia (China)	59

Note: Tributaries are excluded.

Total discharge all the world's rivers:  
 $37,000 \text{ km}^3/\text{yr} = 3.7 \cdot 10^{16} \text{ liters/yr}$



The result of Joly's calculations was an age for the Earth that differed little from Kelvin's:

The quantity of sodium now in the sea, and the annual rate of its supply by the rivers, lead, it will be seen, to the deduction that the age of the earth is  $99 \times 10^6$  years. Certain deductions from this are, it will be shown, warranted, so that the final result of this paper will be to show that the probable age is about  $89 \times 10^6$  years. Also, that this is probably a major limit, and that considerable departure from uniformity of activities could hardly amend it to less than  $80 \times 10^6$  years (Joly, 1899: 249).

Joly considered other possible sources of error, including the Na permanently removed from the system as salt deposits and the possible violations of his assumption of a constant rate of Na influx, but concluded that, taken as a whole, they probably were insignificant. He concluded

We think that it is at least justifiable to claim that our present knowledge of solvent denudation of the earth's surface points to a period of between eighty and ninety millions of years having elapsed since water condensed upon the earth, and rain and rivers and the actions continually progressing in the soils began to supply ocean with materials dissolved from the rocks (Joly 1899: 287).

BILLIONS OF YEARS AFTER UNIVERSE ORIGIN

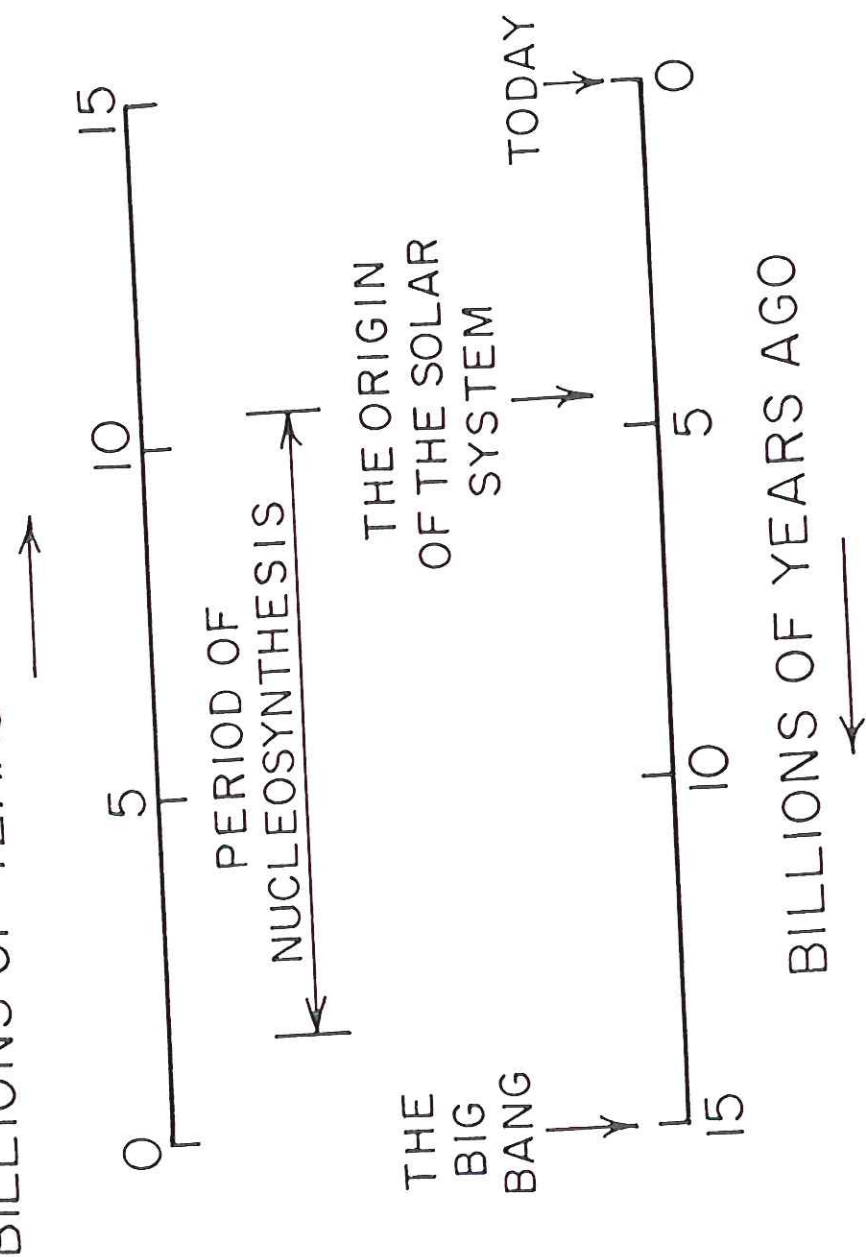
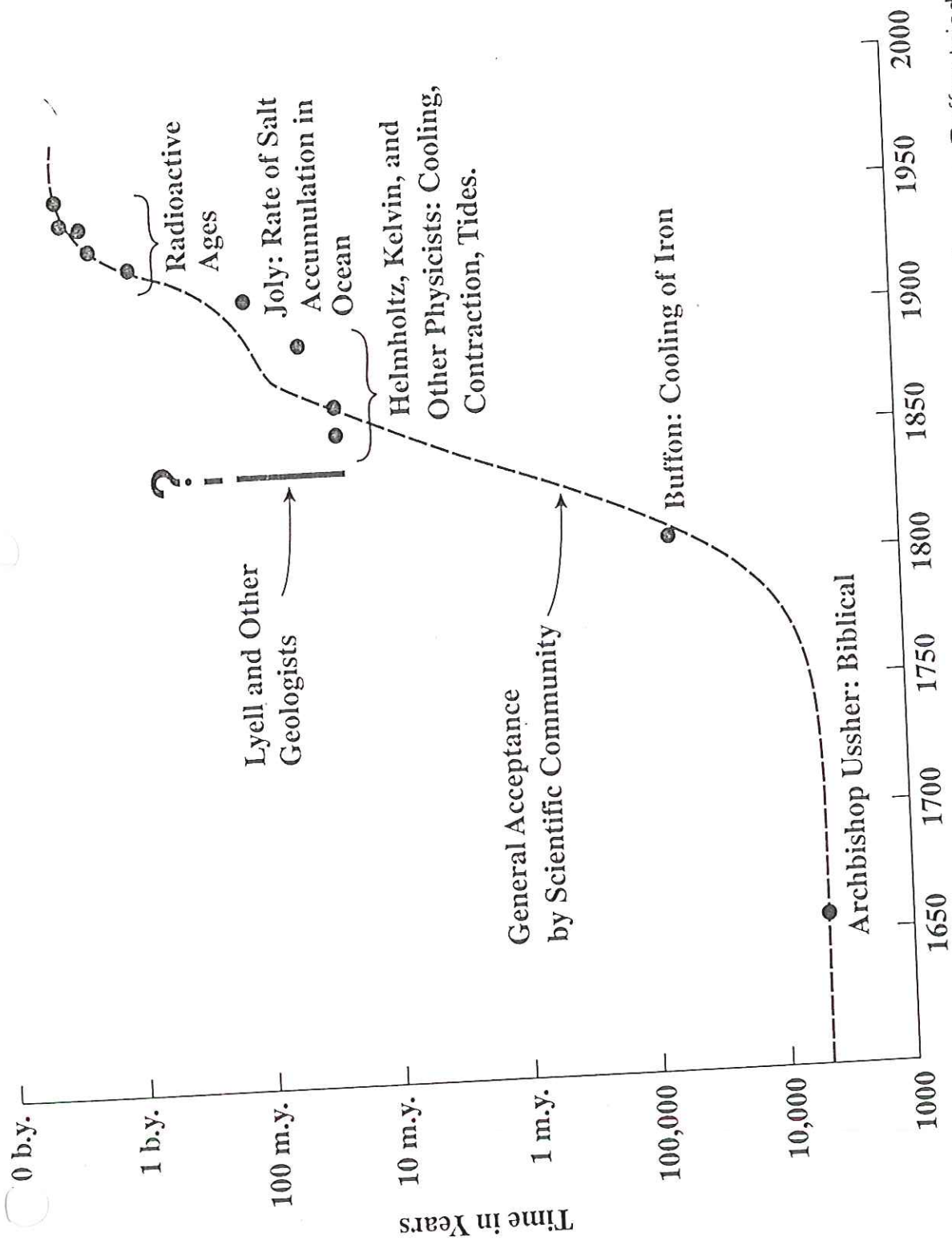


Figure 4-7. Summary of the chronology of universe events: The period of nucleosynthesis refers to the time interval over which the elements heavier than hydrogen and helium that are found in our solar system were produced. For the galaxy as a whole the period of nucleosynthesis extends right up to the present. The matter in the solar system was isolated from the galaxy 4.6 billion years ago.





**Figure 8.1.** Age of Earth as estimated by various techniques since the Renaissance. Buffon tried to use rate of cooling of iron from a molten state to estimate the age of Earth, but the omission of as-yet undiscovered radioactive heating of Earth's interior (chapters 9 and 11) seriously shortened his cooling time. Joly worked out how long it would take to bring the oceans up to their current salinity based on the rate at which rivers carry salt to the sea; he ignored the precipitation of salt out of the ocean water into seafloor sediments. From Press and Siever (1978) by permission of W.H. Freeman and Company.

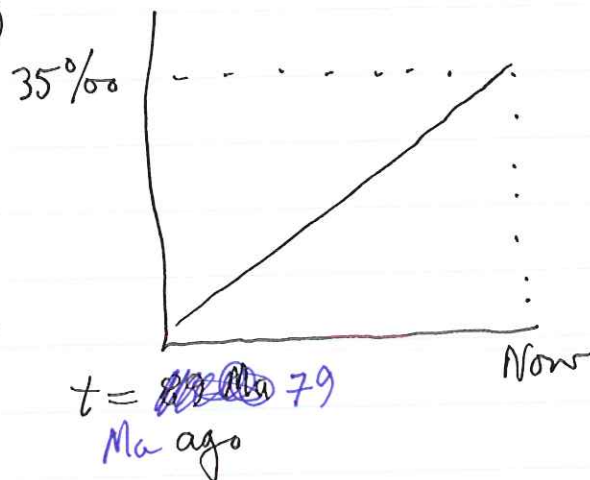
Actual age of  $\oplus$  (Broecker Fig 4-7)  
is 4.7 billion years

So Joly is off by a factor of 50  
(Lunine Fig. 8.1 again)

What is wrong — Joly neglected  
salt deposition → see his quote — he  
actually considered it but regarded it as small  
Evaporite deposition commonly occurs  
in silled basins in warm arid  
climates

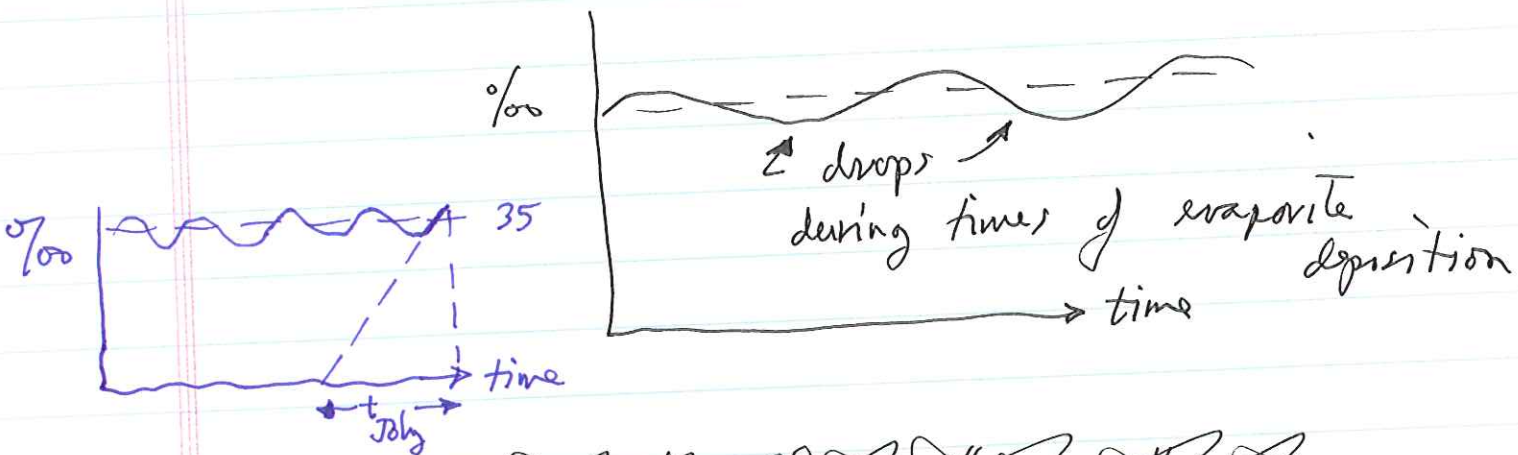
Evaporation of  $H_2O$  leaves behind  
the salts  $NaCl$ ,  $CaSO_4$   
halite, gypsum

Joly imagined that the salinity  
of the ocean was constantly  
increasing





In fact the long-term average salinity is essentially constant in time



~~ohy is the age of the drops mainly in residence time of the salt in the oceans~~

For example, entire Gulf of Mexico underlain by ~ 1 km thick layer of evaporite (salt)

Area of Gulf  $1.6 \cdot 10^6 \text{ km}^2$  (World Almanac)

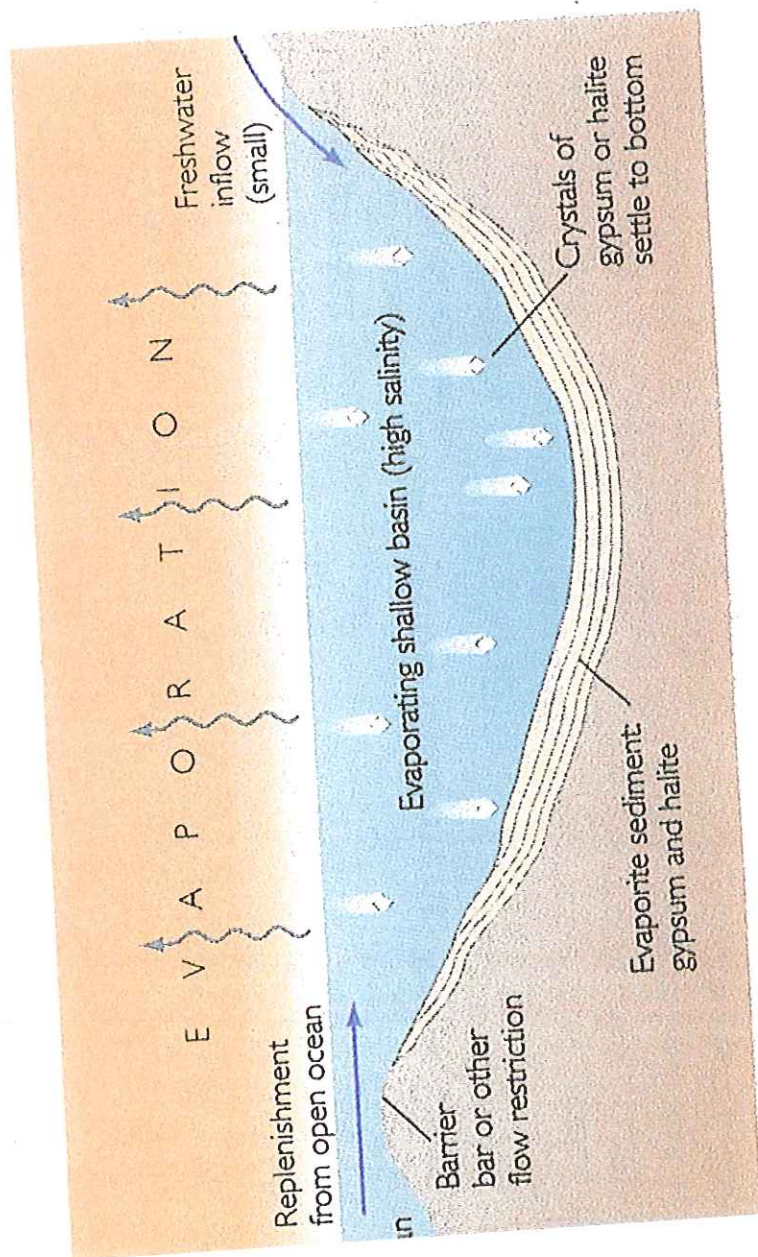
$$\begin{aligned}
 &1.6 \cdot 10^6 \text{ km}^3 \text{ salt} \\
 &= 1.6 \cdot 10^{15} \text{ m}^3 \\
 &= 3.5 \cdot 10^{18} \text{ kg salt}
 \end{aligned}$$

$\rho_{\text{NaCl}} = 2200 \text{ kg/m}^3$

$$\frac{3.5 \cdot 10^{21} \text{ g}}{(24 + 35.5) \text{ g/mole}} = 6 \cdot 10^{19} \text{ moles}$$

1/10 as much Na<sup>+</sup> in this one evaporite deposit as in all oceans

**FIGURE 7.21** When seawater evaporates in a shallow basin with a restricted connection to the open ocean, gypsum and halite form as evaporite sediments. Evaporation removes much more water than the fresh water flowing in can replace. As the evaporating basin gets appreciably more saline than the water of the open sea, gypsum is precipitated. A further increase in salinity leads to the crystallization of halite.





*Evolution of the northern Gulf of Mexico*

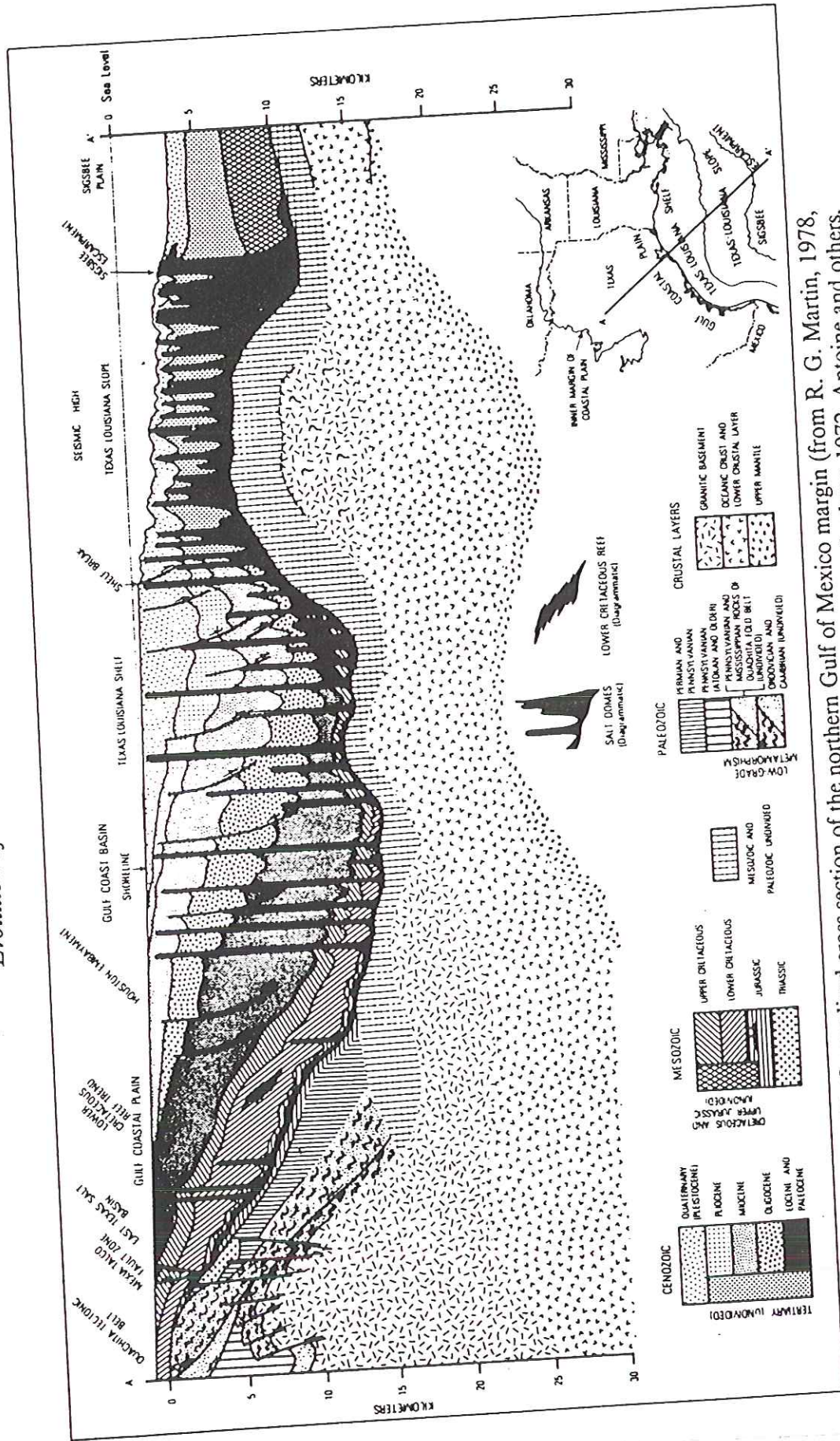


Figure 12. Generalized cross section of the northern Gulf of Mexico margin (from R. G. Martin, 1978, modified from earlier interpretations of Leher, 1969; Dorman and others, 1972; Antoine and others, 1974; and Martin and Case, 1975).

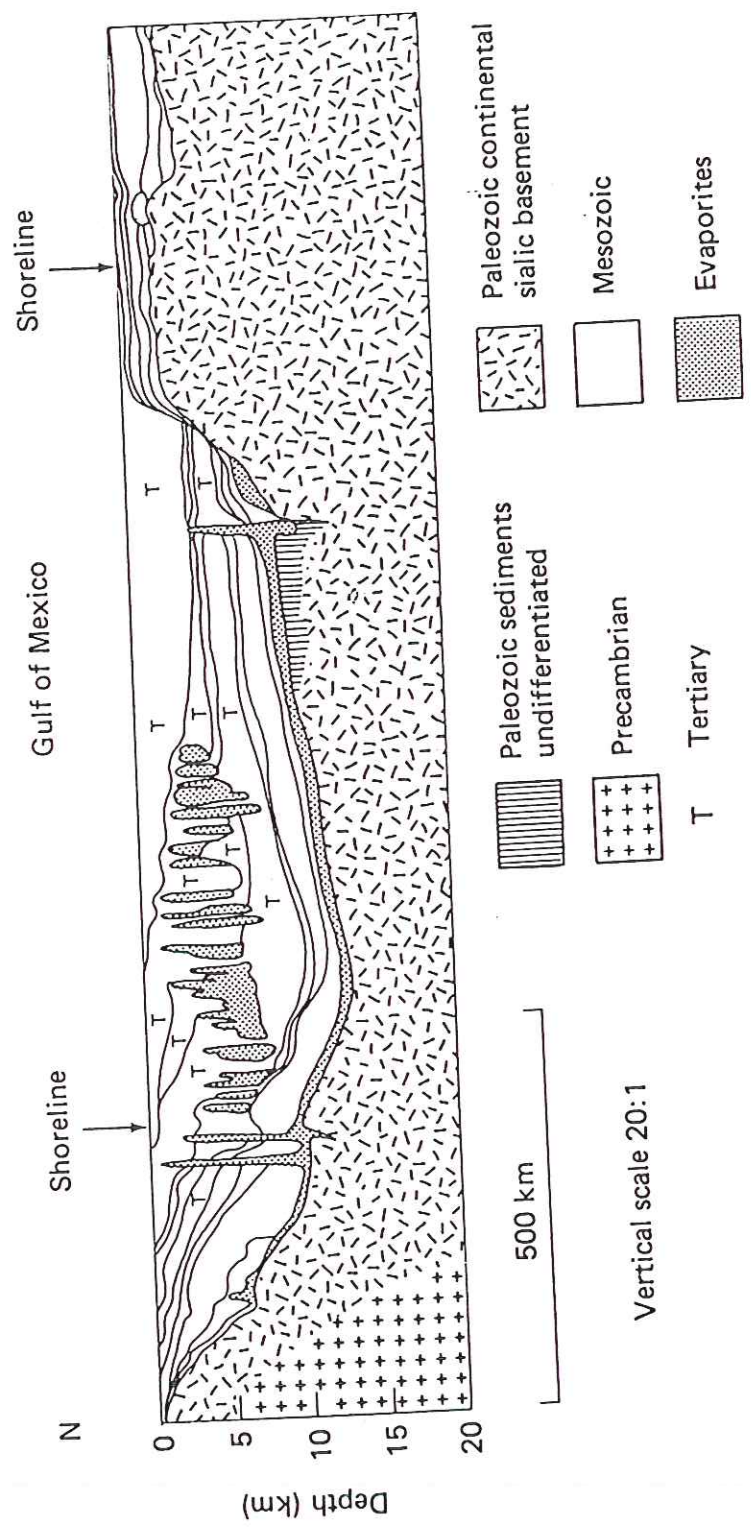


Figure 11-18 Geological cross section for the Gulf of Mexico. Vertical exaggeration is 20:1. (After Bally, 1979)



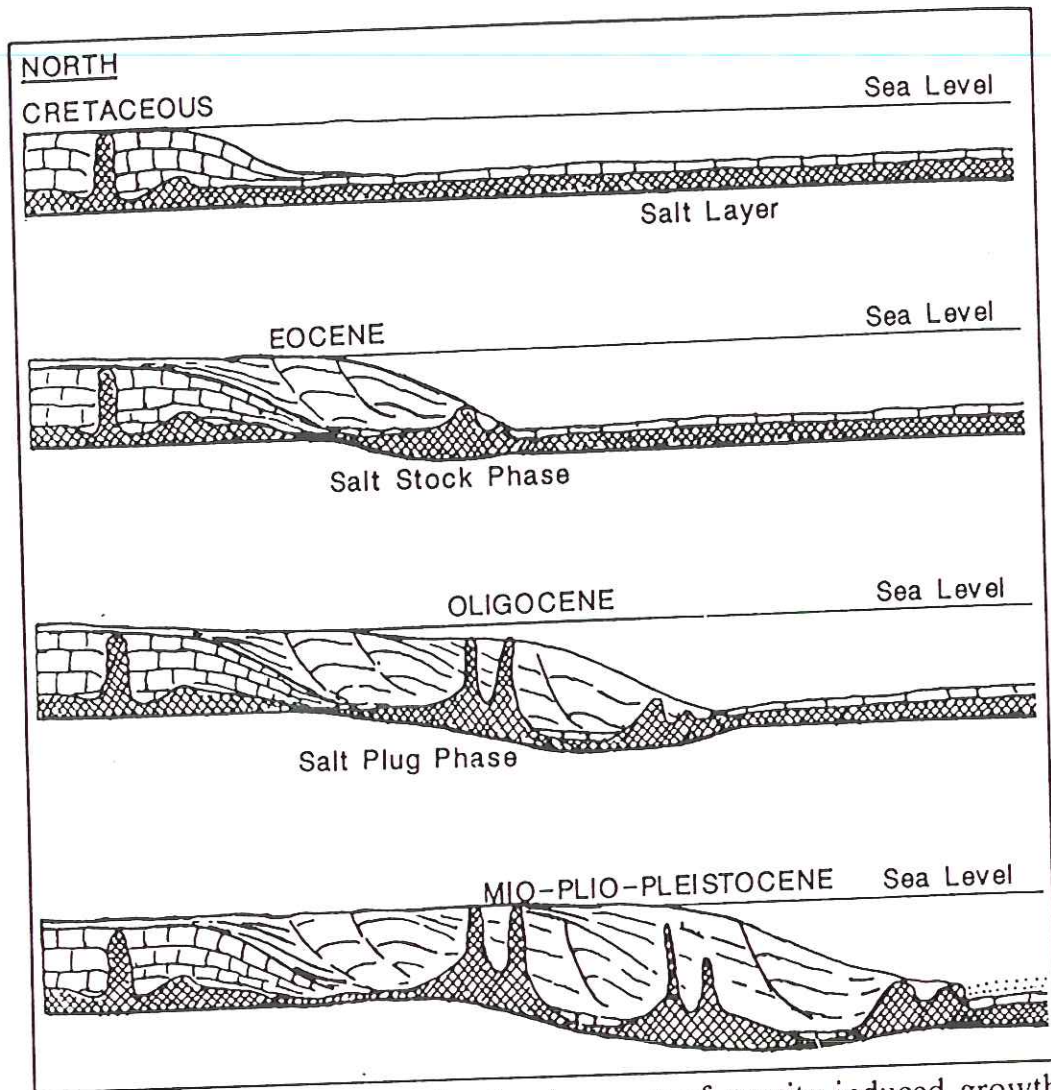


Figure 10. Hypothesis for the development of gravity-induced growth faults on a deep mobile salt layer, northern Gulf of Mexico (from Wilhelm and Ewing, 1972).

### SEDIMENT LOADING

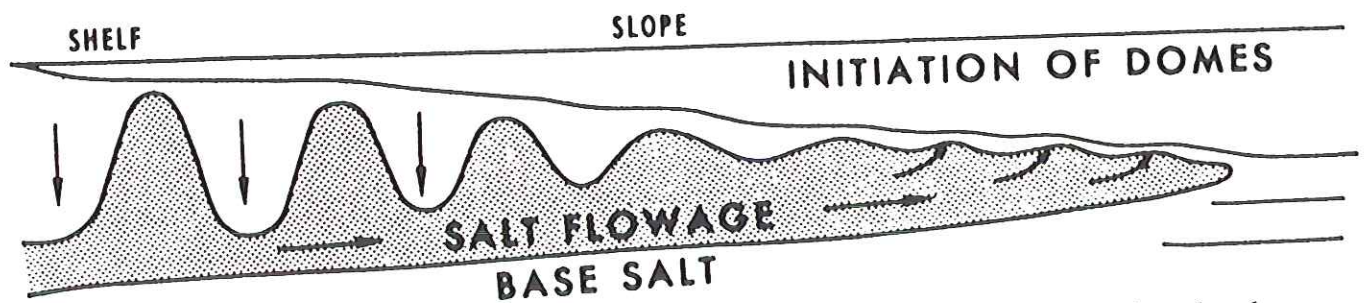


Figure 13. Diagrammatic representation of the initiation of salt dome growth on the continental slope as a result of sediment loading on the shelf-upper slope (from Humphris, 1978, 1979).



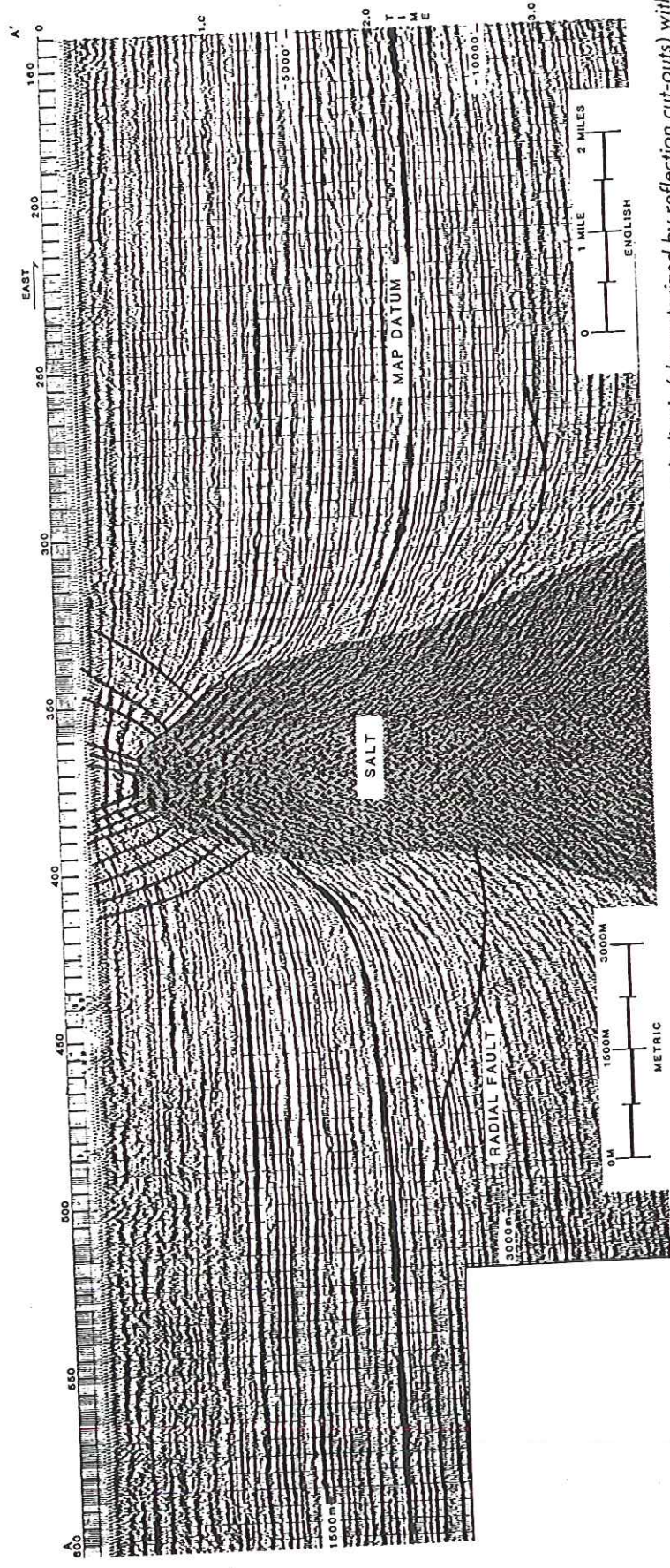


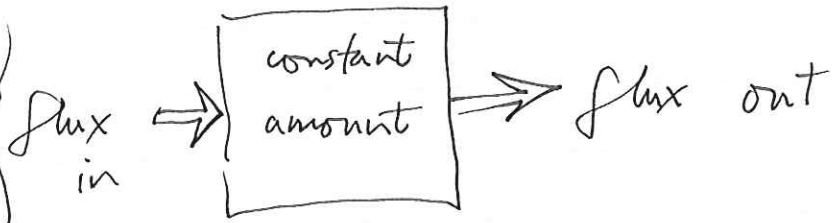
Fig. 8-31 (Sunwall et al., 1983)—Seismic line from offshore Louisiana, U. S. Gulf Coast, showing young salt diapir (characterized by reflection cut-outs) with superjacent normal faults and lower radial faults which strike parallel to seismic line. Salt configuration based on well control, gravity, and reflection and refraction seismic. Note that stratigraphic thinning toward diapir begins early (below 2 sec), but secondary rim synclines are not developed. Permission to publish by American Association of Petroleum Geologists.



Tolys "age of the  $\oplus$ " now viewed as a determination of the residence time of an  $\text{Na}^+$  atom in the oceanic reservoir

Steady-state box model :

Blackboard discussion before showing box model transparency



Example - students at Princeton

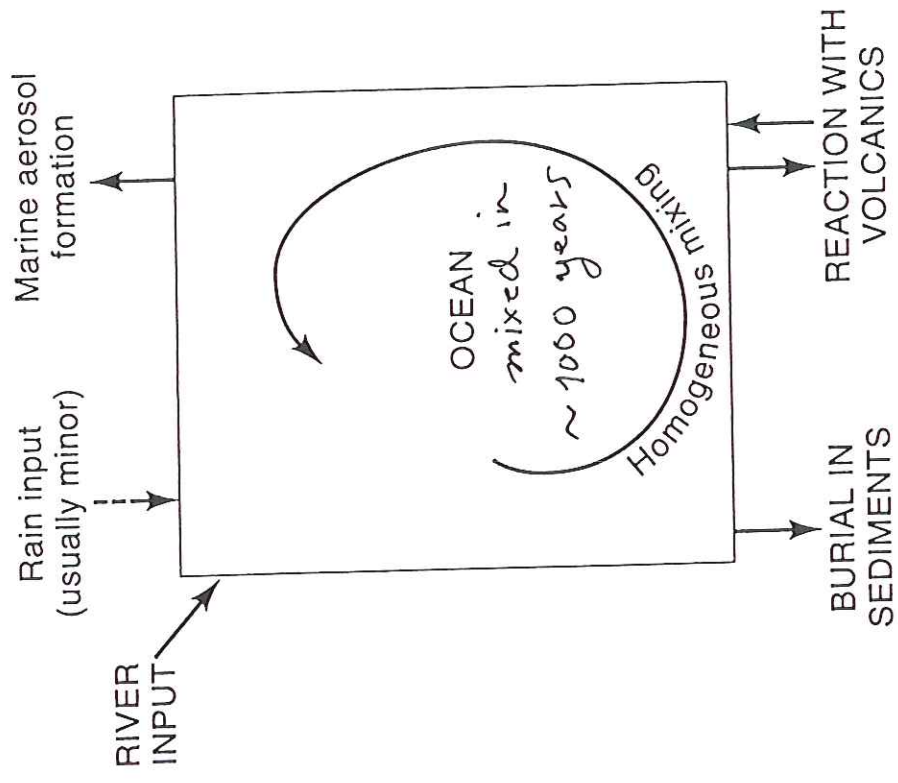
- observe # Princeton students = 4400
- observe # incoming freshmen = 1100
- or # caps & gowns on Nassau Green on Commencement day = 1100

Cheryl Subarban  
dismissing  
freshmen  
at  
beginning  
of  
Princeton  
Week

Residence time =  $\frac{4400 \text{ students}}{1100 \text{ students/yr}} = 4 \text{ yrs}$

Show Berner Tables ~~8.1~~ and J.6 again.

Can be used to determine the residence



**Figure 8.3.** Simple box model appropriate for conservative elements in seawater. Note that, compared with the Sillén model of Figure 8.2, fluxes enter and leave the box, and the atmosphere and sediments are considered to be outside the box. Also note that, in contrast to lakes, there is no outlet, so dissolved materials carried in by rivers can be removed only by sea-air transfer (marine aerosol formation), burial in sediments, or reaction with volcanics.



**TABLE 8.3** Replacement Time with Respect to River Addition,  $\tau_r$ , for Some Major and Minor Dissolved Species in Seawater

Component	Concentration ( $\mu\text{M}$ )		$\tau_r^a$ (1,000 yr)
	River Water	Seawater	
Cl	230	558,000	87,000
Na <sup>+</sup>	315	479,000	55,000
Mg <sup>++</sup>	150	54,300	13,000
SO <sub>4</sub> <sup>--</sup>	120	28,900	8,700
Ca <sup>++</sup>	367	10,500	1,000
K <sup>+</sup>	36	10,400	10,000
HCO <sub>3</sub> <sup>-</sup>	870	2,000	83
H <sub>4</sub> SiO <sub>4</sub>	170	100	21
NO <sub>3</sub> <sup>-</sup>	10	20	72
Orthophosphate	1.8	2	40

*only not 79. 10<sup>6</sup>?*  
*uses flux value including modern pollution*

<sup>a</sup>  $\tau_r = ([\text{SW}]/[\text{RW}]) \tau_w$ , where  $\tau_w$  = replacement (residence) time of H<sub>2</sub>O = 36,000 yr; RW = river water; SW = seawater, and [ ] = concentration in  $\mu\text{moles per liter} = \mu\text{M}$ .

times of several species

Table 8.3 gives these — uses the modern (polluted) riverine inputs

$$\frac{\text{Na}^+ \text{ (polluted)}}{\text{Na}^+ \text{ (natural)}} = \frac{8.4}{6.5} = 1.3 \text{ times}$$

as great a flux

$$\frac{79 \text{ my}}{1.3} = 61 \text{ myr} - \text{ ~~61 myr~~ Berner gives 35 myr}$$

A major aim of geochemistry is to balance the books on a given element. Briefly consider ~~the~~ four

- Cl<sup>-</sup> chlorine — all processes well identified — not incorporated in silicate minerals — only in evaporites NaCl

Table 8.13 one Tg (teragram) = 10<sup>12</sup> g  
= 10<sup>9</sup> kg

net air-sea transfer

sea breeze picks up aerosols

↑ can smell the ocean

near the shore



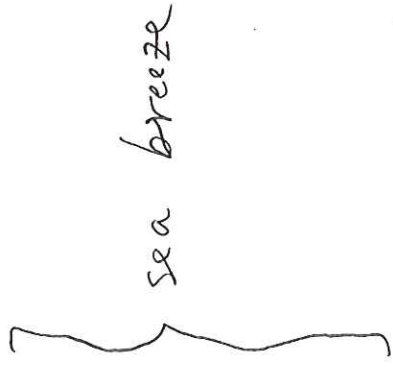
**TABLE 8.13** The Oceanic Chloride Budget (Rates in Tg Cl<sup>-</sup>/yr)

Present-Day Budget	
Inputs	Outputs
Rivers (natural)	Net sea-air transfer
Rivers (pollution)	Pore-water burial
Total	Total
215	40
93	25
<u>308</u>	<u>65</u>
Long-Term (Balanced) Budget	
Inputs	Outputs
Rivers	NaCl evaporative deposition
	Net sea-air transfer
	Pore-water burial
	Total
215	163
	40
	12
	<u>215</u>

Note: Tg = 10<sup>12</sup> g. Replacement time for Cl<sup>-</sup> is 87 million years.

**TABLE 8.12** Rates of Addition via Rivers of Major Elements to the ocean (as Dissolved Species) and Rates of Net Loss from the Ocean by Transfer of Sea Salt to the Continents via the Atmosphere

Species	Rate of Addition from Rivers <sup>a</sup> (Tg/yr)	Rate of Net Sea Salt Loss to Atmosphere (Tg/Yr)
Cl <sup>-</sup>	308	40
Na <sup>+</sup>	269	21
SO <sub>4</sub> -S	143	4
Mg <sup>++</sup>	137	3
K <sup>+</sup>	52	1
Ca <sup>++</sup>	550	0.5
HCO <sub>3</sub> <sup>-</sup>	1980	—
H <sub>4</sub> SiO <sub>4</sub> -Si	180	—



Note: Tg = 10<sup>12</sup> g.

<sup>a</sup> Based on river water input of 37,400 km<sup>3</sup>/yr: includes pollution

Sources: River-water data from Meybeck 1979; cyclic salt data from Chapter 5.



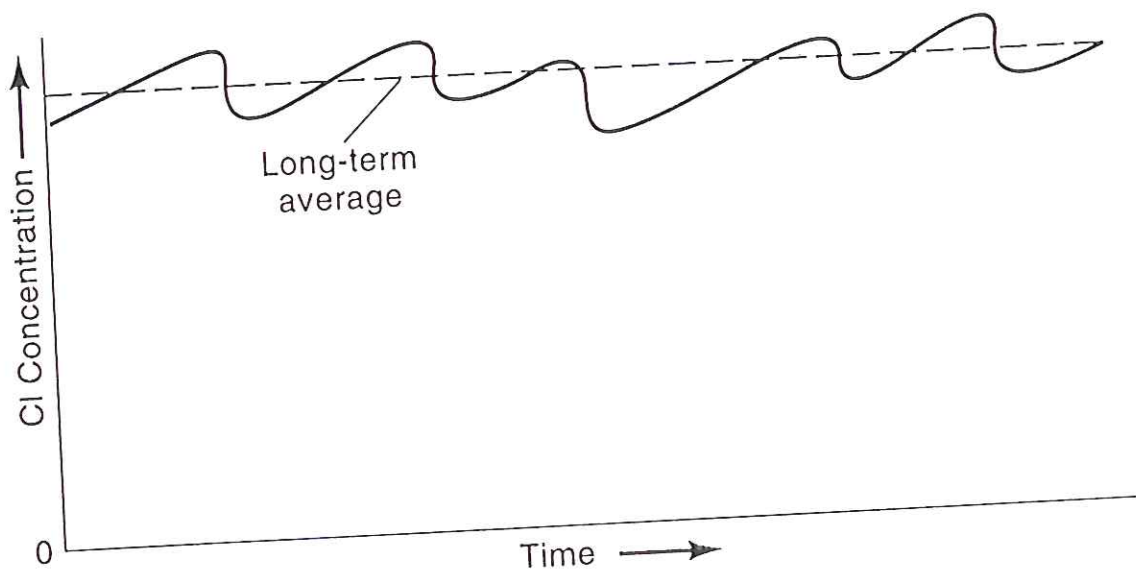
**TABLE 8.13** The Oceanic Chloride Budget (Rates in Tg Cl<sup>-</sup>/yr)

Present-Day Budget			
Inputs		Outputs	
Rivers (natural)	215	Net sea-air transfer	40
Rivers (pollution)	93	Pore-water burial	25
Total	308	Total	65

Long-Term (Balanced) Budget			
Inputs		Outputs	
Rivers	215	NaCl evaporative deposition	163
		Net sea-air transfer	40
		Pore-water burial	12
		Total	215

Note: Tg = 10<sup>12</sup> g. Replacement time for Cl<sup>-</sup> is 87 million years.



**Figure 8.12.** Schematic representation of change of chloride (Cl<sup>-</sup>) concentration in seawater with geologic time. Sudden drops are due to the rapid precipitation of NaCl in evaporite basins.

**TABLE 4.7** Origin of Major Components of Groundwater (Major Processes Only)

Component	Origin
Na <sup>+</sup>	NaCl dissolution (some pollutive) <sup>a</sup> Plagioclase weathering Rainwater addition
K <sup>+</sup>	Biotite weathering K-feldspar weathering Biomass decreases Dissolution of trapped aerosols
Mg <sup>++</sup>	Amphibole and pyroxene weathering Biotite (and chlorite) weathering Dolomite weathering Olivine weathering Rainwater addition
Ca <sup>++</sup>	Calcite weathering Plagioclase weathering Dolomite weathering Dissolution of trapped aerosols Biomass decreases
HCO <sub>3</sub> <sup>-</sup>	Calcite and dolomite weathering Silicate weathering
SO <sub>4</sub> <sup>--</sup>	Pyrite weathering (some pollutive) <sup>a</sup> CaSO <sub>4</sub> dissolution Rainwater addition
Cl <sup>-</sup>	NaCl dissolution (some pollutive) <sup>a</sup> Rainwater addition
H <sub>4</sub> SiO <sub>4</sub>	Silicate weathering

*Note:* Order presented is approximate order of decreasing importance.



**TABLE 5.11** Sources of Major Elements in World River Water (in Percent of Actual Concentrations)

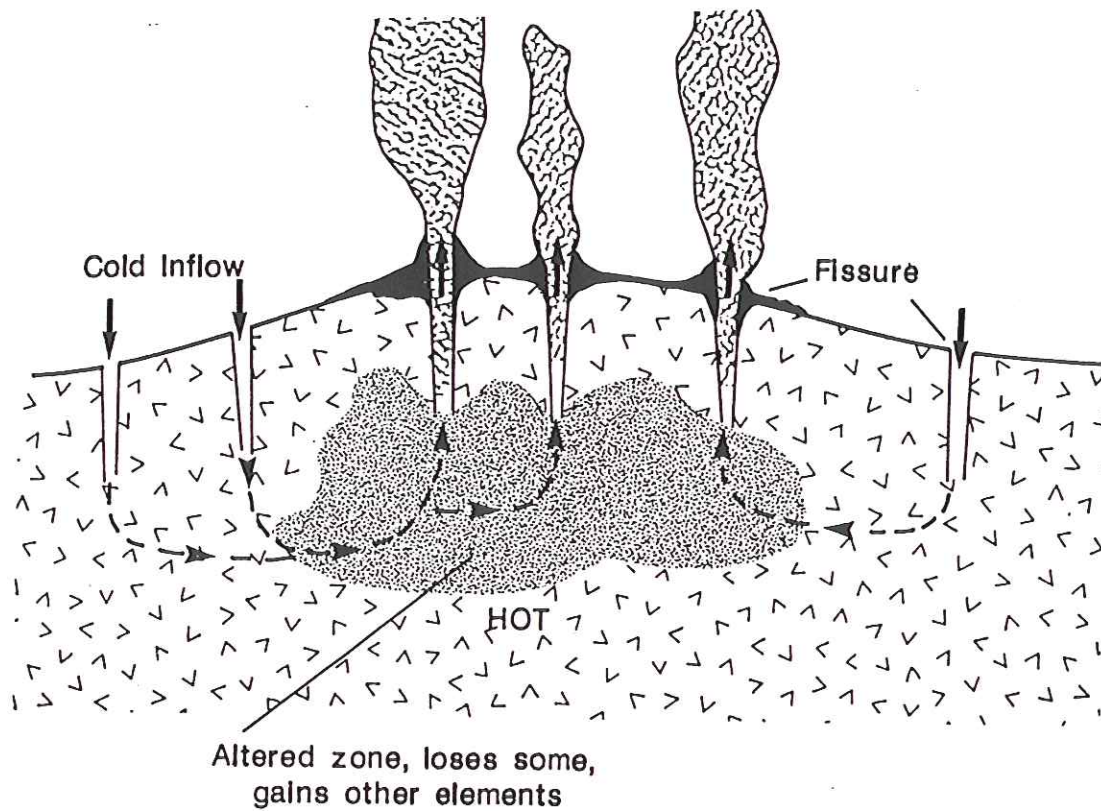
Element	Weathering				
	Atmos. Cyclic Salt	Carbonates	Silicates	Evaporites <sup>a</sup>	Pollution <sup>b</sup>
Ca <sup>++</sup>	0.1	65	18	8	9
HCO <sub>3</sub> <sup>-</sup>	<<1	61 <sup>c</sup>	37 <sup>c</sup>	0	2
Na <sup>+</sup>	8	0	22	42	28
Cl <sup>-</sup>	13	0	0	57	30
SO <sub>4</sub> <sup>--</sup>	2 <sup>d</sup>	0	0	22 <sup>d</sup>	54
Mg <sup>++</sup>	2	36	54	<<1	8
K <sup>+</sup>	1	0	87	5	7
H <sub>4</sub> SiO <sub>4</sub>	<<1	0	99+	0	0

<sup>a</sup> Also includes NaCl from shales and thermal springs.

<sup>b</sup> Values taken from Meybeck (1979) except sulfate, which is based on calculation given in the text.

<sup>c</sup> For carbonates, 34% from calcite and dolomite and 27% from soil CO<sub>2</sub>; for silicates, all 37% from soil CO<sub>2</sub>; thus, total HCO<sub>3</sub><sup>-</sup> from soil (atmospheric) CO<sub>2</sub> = 64% (See also Table 5.13).

<sup>d</sup> Other sources of river SO<sub>4</sub><sup>--</sup>; natural biogenic emissions to atmosphere delivered to land in rain, 3%; volcanism, 8%; pyrite weathering, 11%.



**TABLE 8.8** Removal or Addition Fluxes of Some Major Seawater Constituents as a Result of Basalt-Seawater Reaction Near Mid-ocean Ridges

Constituent	Flux (Tg/yr)		
	Edmond et al. (1979)	Thompson (1983)	Present Study
Mg <sup>++</sup>	-187	-60	-119
Ca <sup>++</sup>	140	73	191
K <sup>+</sup>	51	-27	53
SO <sub>4</sub> <sup>--</sup> as S	-120	—	— <sup>a</sup>
H <sub>4</sub> SiO <sub>4</sub> as Si	90	82	56

*Note:* Removal values shown as negative numbers. Tg = 10<sup>12</sup>g.

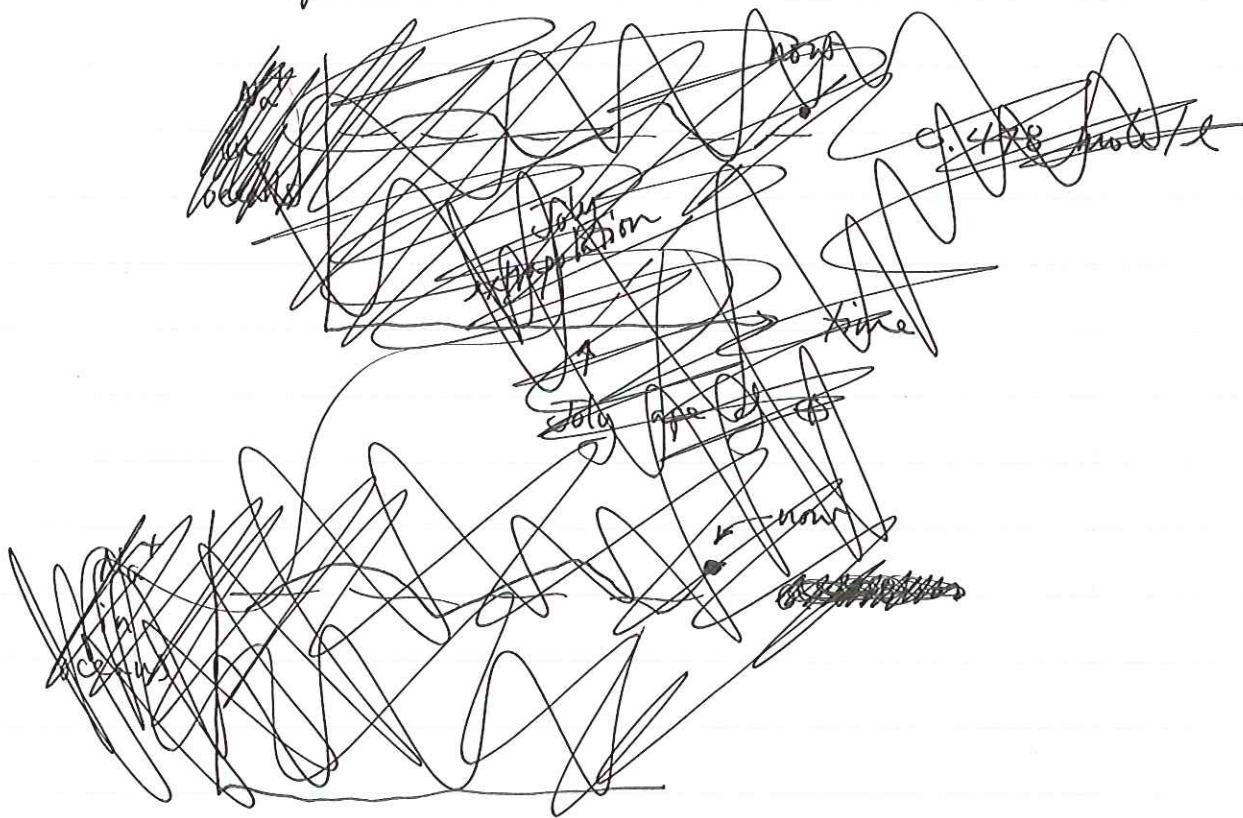
<sup>a</sup> Less than 10% of Edmond et al. (1979) value



burial in fore sediments

Major imbalance at present time

- NaCl road salt 93/308
- time of reduced evaporite deposition



why — sea level particularly high now  
interglacial

•  $\text{Na}^+$  — Toly's element

more sources & sinks

Na in feldspars —

Na in clays

ocean basalt

plagioclase  $\text{NaAlSi}_3\text{O}_8$

Table 4.7

Table 8.14 — present day out of balance as for  $\text{Na}^+$

long-term  $\text{NaCl}$  deposition fixed using  $\text{Na}^+$

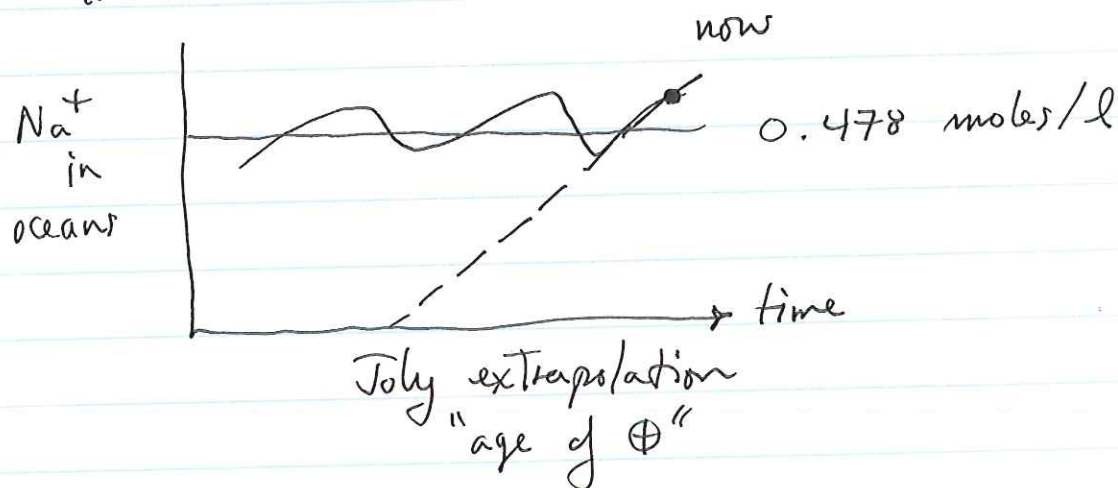
$$163 \times \frac{24}{35.5} = 106$$

cation exchange — clay minerals in estuaries

basalt-seawater reaction — black smokers  
removes  $\text{Na}^+$  from seawater

~~For both  $\text{Na}^+$  and  $\text{Cl}^-$  the picture is the same~~

~~For both  $\text{Na}^+$  and  $\text{Cl}^-$  the picture is the same~~



Consider two more cations



**TABLE 8.14** The Oceanic Sodium Budget (Rates in Tg Na<sup>+</sup>/yr)

Present-Day Budget	
Inputs	Outputs
Rivers (natural)	Cation exchange
Rivers (pollution)	Net sea-air transfer
	Pore-water burial
Total	Total
193	42
76	21
—	16
269	79
Long-Term Budget	
Inputs	Outputs
Rivers	NaCl deposition
	Net sea-air transfer
	Cation exchange
	Pore-water burial
	Basalt-seawater reaction
	Total
193	106
	21
	21
	8
	37
	193

*Note:* Tg = 10<sup>12</sup> g. Replacement time for Na<sup>+</sup> is 55 million years.

• Mg++ magnesium

Table 8.16

Before discovery of MOR hot springs  
 = a serious imbalance problem

Now known that almost all  
 Mg++ added by rivers  
 (119 out of 137 Tg/yr)  
 is removed by reactions with  
 volcanic rock at ridges

• Ca++ calcium

Ca++ and Mg++ same column  
 periodic table — Table 8.3

	conc in river water	both $\mu\text{M}$	conc in seawater
Mg++	150		54,000
Ca++	367		10,500

Much more Ca++ coming in rivers  
 but much less present in oceans

Reason — Ca++ is preferentially  
 incorporated into corals and the shells  
 of marine plankton — Fig. 8.9





$\text{CaCO}_3$  shells

— — — — — grow in photic zone  
top 100 m

Shell growth removes  $\text{Ca}^{++}$  from seawater  
die & sink to bottom.

Berner's budget is balanced to within  
the uncertainties

Most input from rivers

$\text{Ca}^{++}$  also added to ocean by clay  
basalt-seawater reactions (opposite  
of  $\text{Na}^+$  and  $\text{Mg}^{++}$ )

Most output as  $\text{CaCO}_3$  limestone

- shallow water — shelves — coral reefs, etc.
- deep water — plankton

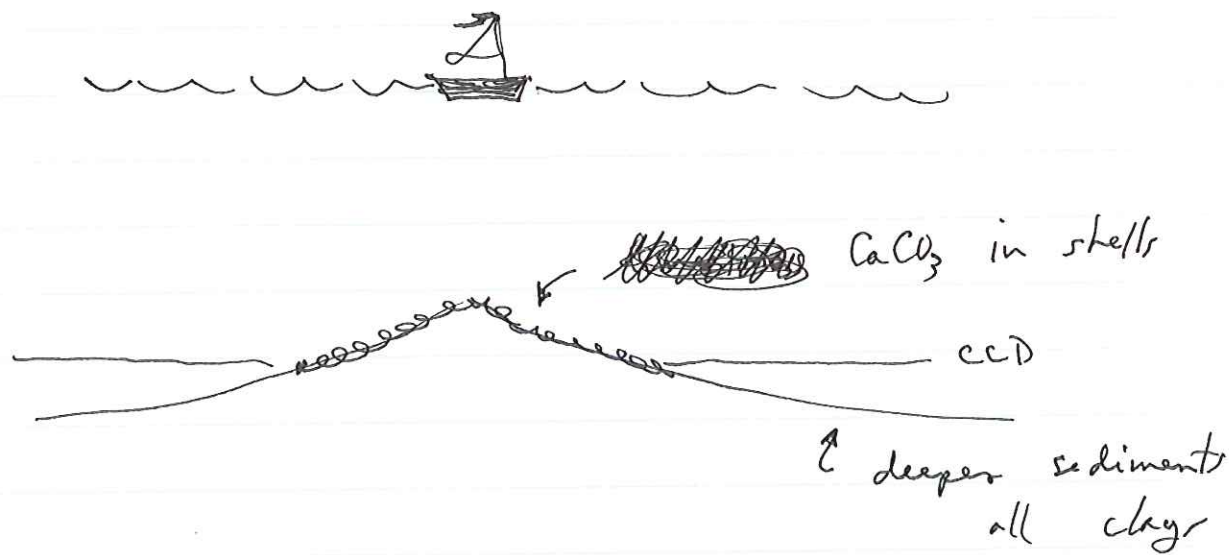
Below  $\sim 4$  km depth the oceans are undersaturated with respect to  $\text{CaCO}_3$

The solubility increases with pressure

As a result deep-sea sediments below  $\sim 4$  km depth dissolve back into water column — calcite compensation depth

See map of Atlantic Fig. 8.13

Think of mid-ocean ridges like a snow-capped peaks



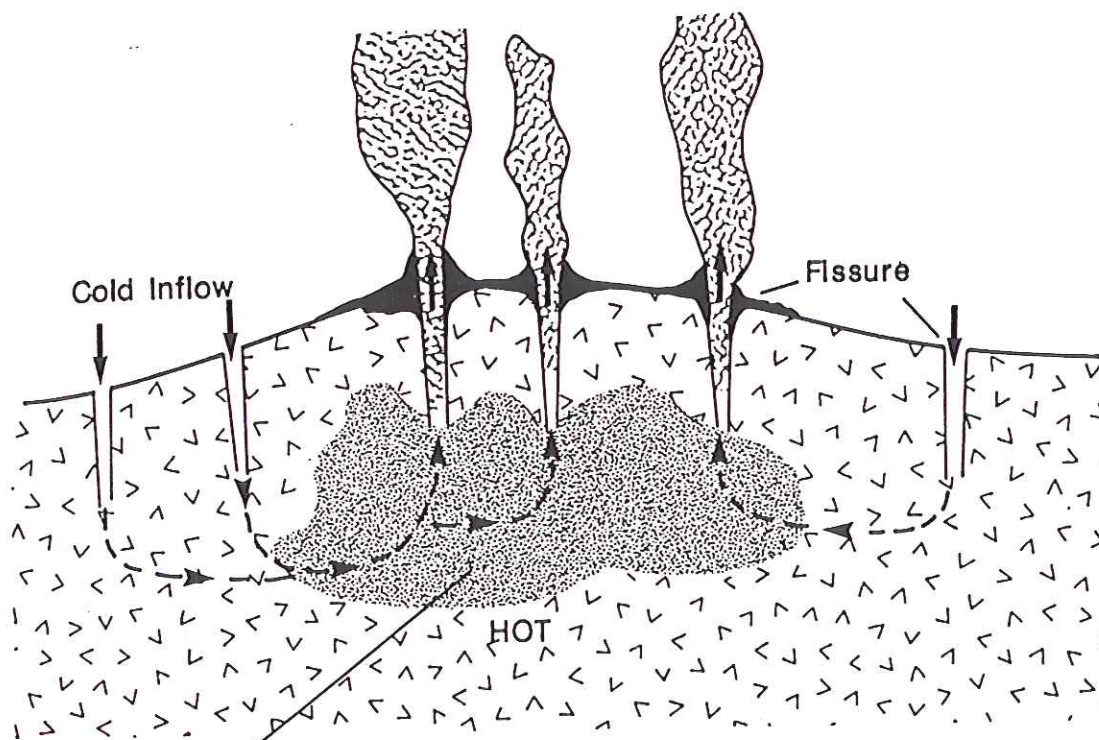


**TABLE 8.16** The Oceanic Magnesium Budget (Rates in Tg Mg<sup>++</sup>/yr)

(Balanced) Budget for Past 100 Million Years

Inputs		Outputs	
Rivers	137	Volcanic-seawater reaction	119
		In biogenic CaCO <sub>3</sub>	15
		Net sea-air transfer	3
		Total	137

Note: Tg = 10<sup>12</sup> g. Replacement time for Mg<sup>++</sup> is 13 million years.



Altered zone, loses some,  
gains other elements

**TABLE 8.16** The Oceanic Magnesium Budget (Rates in Tg Mg<sup>++</sup>/yr)  
(Balanced) Budget for Past 100 Million Years

Inputs		Outputs	
Rivers	137	Volcanic-seawater reaction	119
		In biogenic CaCO <sub>3</sub>	15
		Net sea-air transfer	3
		Total	137

*Note:* Tg = 10<sup>12</sup> g. Replacement time for Mg<sup>++</sup> is 13 million years.



**TABLE 8.3** Replacement Time with Respect to River Addition,  $\tau_r$ , for Some Major and Minor Dissolved Species in Seawater

Component	Concentration ( $\mu\text{M}$ )		$\tau_r^a$ (1,000 yr)
	River Water	Seawater	
Cl	230	558,000	87,000
Na <sup>+</sup>	315	479,000	55,000
Mg <sup>++</sup>	150	54,300	13,000
SO <sub>4</sub> <sup>--</sup>	120	28,900	8,700
Ca <sup>++</sup>	367	10,500	1,000
K <sup>+</sup>	36	10,400	10,000
HCO <sub>3</sub> <sup>-</sup>	870	2,000	83
H <sub>4</sub> SiO <sub>4</sub>	170	100	21
NO <sub>3</sub> <sup>-</sup>	10	20	72
Orthophosphate	1.8 <sup>b</sup>	2	40

<sup>a</sup>  $\tau_r = ([\text{SW}]/[\text{RW}]) \tau_w$ , where  $\tau_w$  = replacement (residence) time of H<sub>2</sub>O = 36,000 yr; RW = river water; SW = seawater, and [ ] = concentration in  $\mu\text{moles per liter} = \mu\text{M}$ .

<sup>b</sup> Includes input from solubilization of solids.

**TABLE 5.12** Sources of Ca and Mg in World Average River Water

Source	Percent of Total Ca	Percent of Total Mg
Weathering		
Calcite, $\text{CaCO}_3$	52	—
Dolomite, $\text{CaMg}(\text{CO}_3)_2$	13	36
$\text{CaSO}_4$ minerals	8	—
Ca-silicates ( <i>plagioclase</i> )	18	—
Mg-silicates	—	54
Cyclic sea salt	<<1	2
Pollution	9	8
Total	100	100

*Source:* Data for rock sources from Berner et al. 1983. Cyclic sea salt from Table 5.10 and pollution from Meybeck 1979.

**TABLE 5.13** Sources of Rock Weathering-Derived  $\text{HCO}_3^-$  in World Average River Water

Weathering Source	Percent of Total $\text{HCO}_3^-$ from Soil $\text{CO}_2$	Percent of Total $\text{HCO}_3^-$ from Carbonate Minerals
Calcite + Dolomite	27	34
Ca-silicates	13	—
Mg-silicates	15	—
Na-silicates	6	—
K-silicates	3	—
Total	64	34

*Note:* For method of calculation, see text. (An additional 2% of total  $\text{HCO}_3^-$  is added by pollution; see Table 5.11.)



**TABLE 8.18** The Oceanic Calcium Budget (Rates in Tg Ca<sup>++</sup>/yr)

Present-Day Budget			
Inputs		Outputs	
Rivers	550	CaCO <sub>3</sub> deposition:	
Volcanic-seawater reaction	191	Shallow water	520
Cation exchange	37	Deep sea	440
Total	778	Total	960
Budget for Past 25 Million Years			
Inputs		Outputs	
Rivers	550	CaCO <sub>3</sub> deposition:	
Volcanic-seawater reaction	191	Shallow water	240
Cation exchange	19	Deep sea	440
Total	760	Evaporitic CaSO <sub>4</sub> deposition	49
		Total	729

Note: Tg = 10<sup>12</sup> g. Replacement time (rivers only) for Ca<sup>++</sup> is 1 million years.

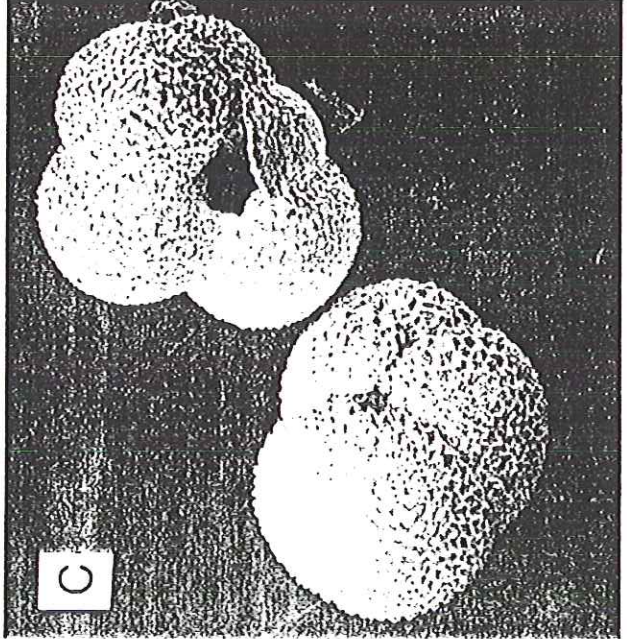
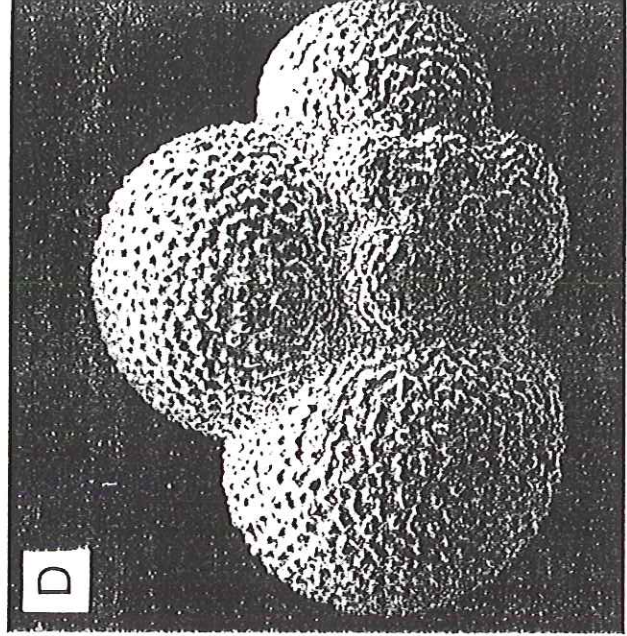
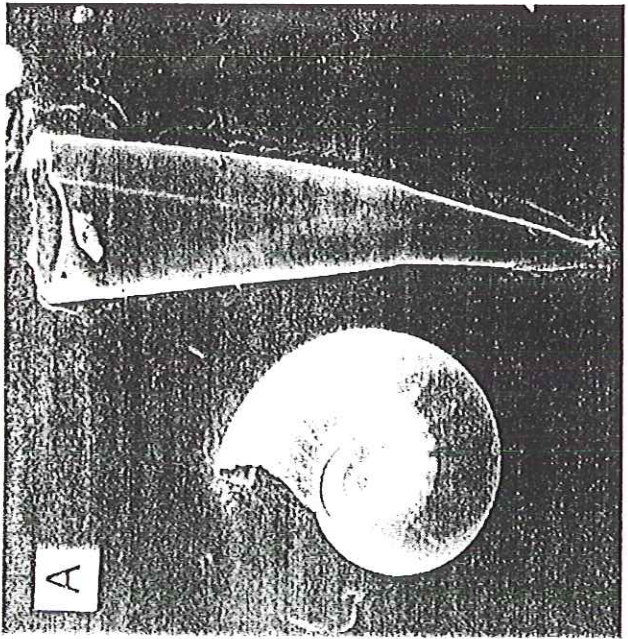
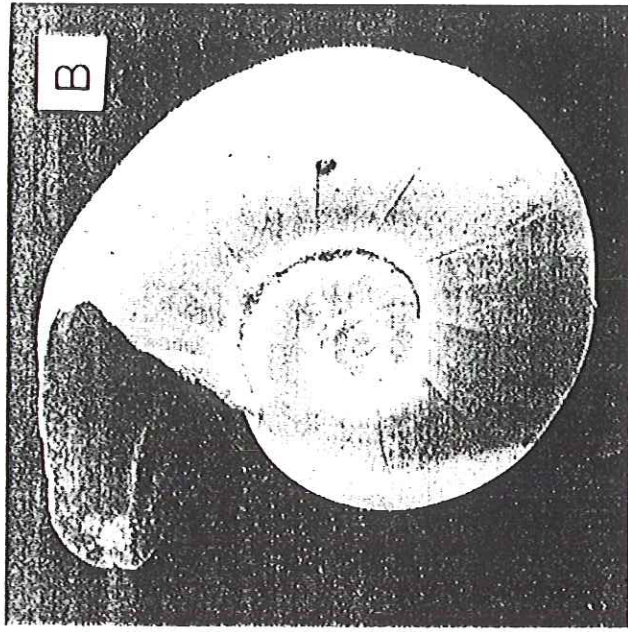
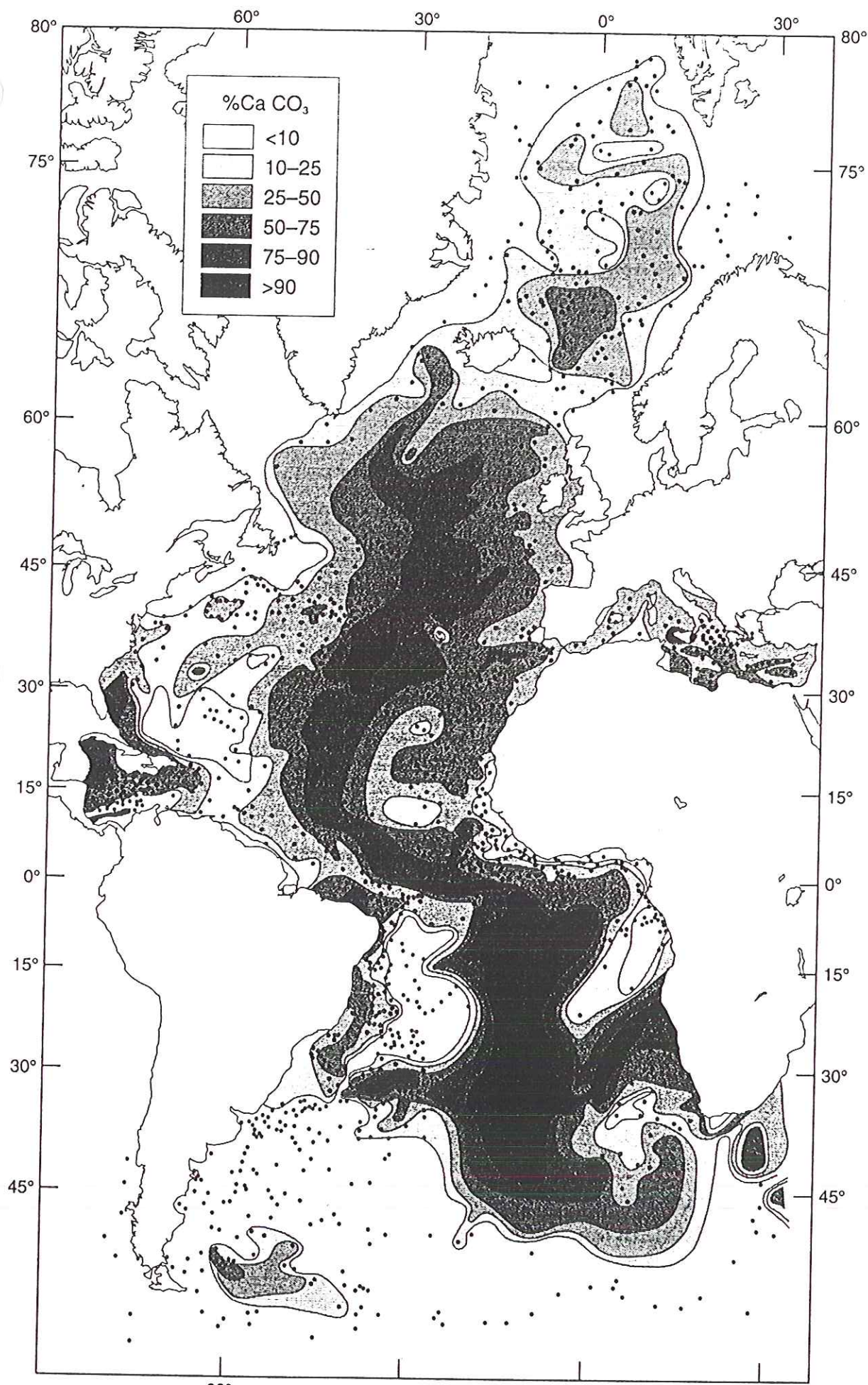


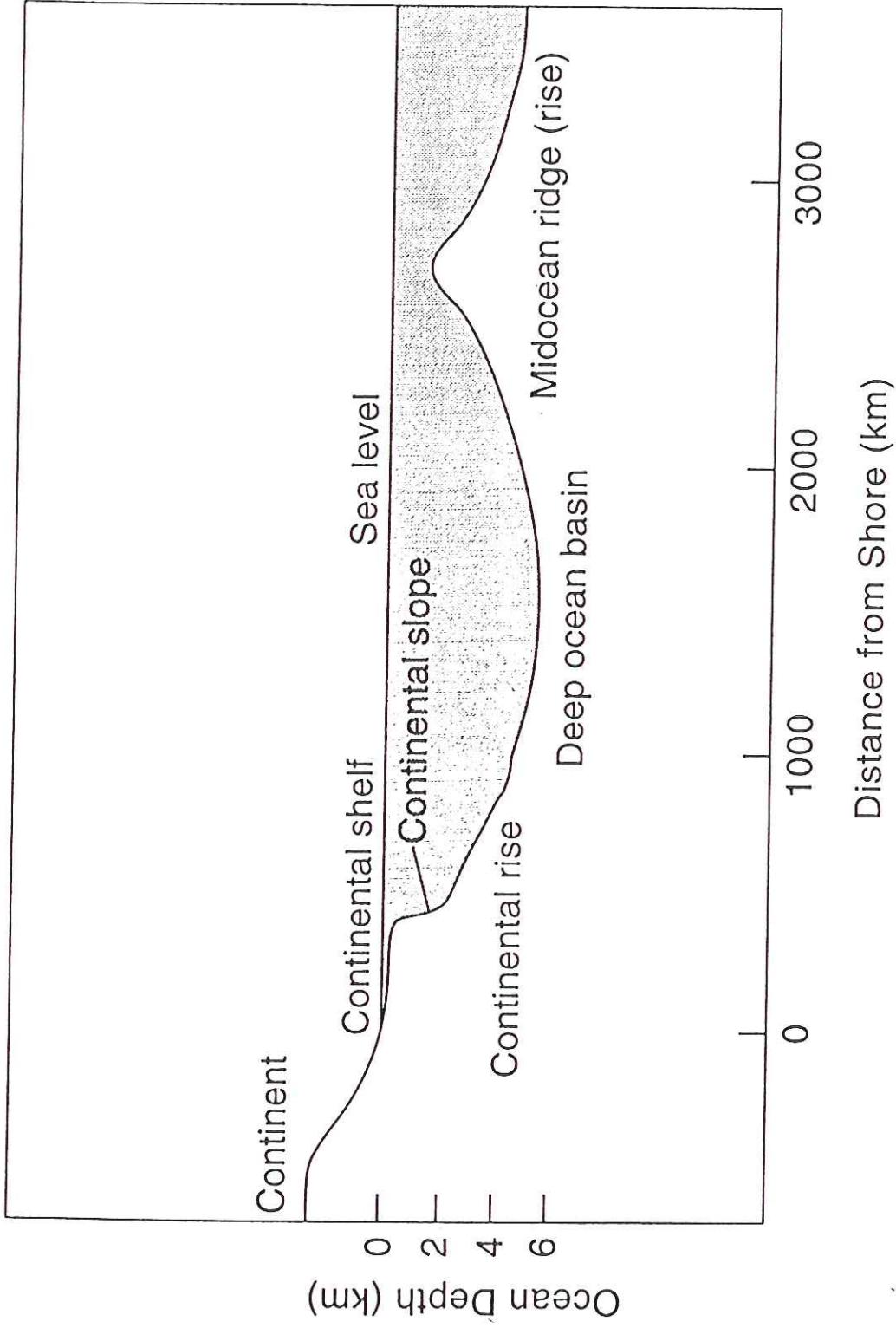
Figure 8.9. Photomicrographs of some planktonic  $\text{CaCO}_3$ -secreting organisms: (a) pteropod shells (aragonite)  $\times 20$ ; (b) pteropod shell (aragonite)  $\times 10$ ; (c) foramin tests (calcite)  $\times 70$ ; (d) foramin test (calcite)  $\times 100$ .





**Figure 8.13.** Distribution of CaCO<sub>3</sub> in deep-sea sediments of the Atlantic Ocean. Note that the highest concentrations are located at the shallowest depths atop the Mid-Atlantic Ridge. (After P. E. Biscaye, V. Kolla, and K. K. Turekian, "Distribution of Calcium Carbonate in Surface Sediments of the Atlantic Ocean." *Journal of Geophysical Research* 81: 2596. Copyright © 1976 by the American Geophysical Union, printed by permission of the publisher.)





**Figure 8.1.** Generalized schematic cross section of the oceans showing major physiographic features. Note large differences in vertical and horizontal scales.

# PERIODIC TABLE OF THE ELEMENTS

<b>Mn</b>	—	Chemical Symbol
25	—	Atomic Number
Manganese	—	Element Name
54.94	—	Atomic Weight

Strong tendency for outermost electrons to be lost to make full outer shell

<b>H</b>	1	Hydrogen	1.01
<b>Li</b>	3	Lithium	6.94
<b>Be</b>	4	Beryllium	9.01
<b>Na</b>	11	Sodium	22.99
<b>Mg</b>	12	Magnesium	24.31

Noble gases: outer shells filled; no tendency to gain or lose electrons

<b>He</b>	2	Helium	4.00
<b>Ne</b>	10	Neon	20.18
<b>Ar</b>	18	Argon	39.95
<b>Kr</b>	36	Krypton	83.80
<b>Xe</b>	54	Xenon	131.30
<b>Rn</b>	86	Radon	(222)

Strong tendency to gain electrons to make full outer shell

<b>F</b>	9	Fluorine	19.00
<b>O</b>	8	Oxygen	16.00
<b>S</b>	16	Sulfur	32.06
<b>Cl</b>	17	Chlorine	35.45
<b>Br</b>	35	Bromine	79.90
<b>I</b>	53	Iodine	126.90
<b>At</b>	85	Astatine	(210)

Tendency to fill outer electron shell by electron sharing and gain or loss of electrons

<b>C</b>	6	Carbon	12.01
<b>N</b>	7	Nitrogen	14.01
<b>O</b>	8	Oxygen	16.00
<b>F</b>	9	Fluorine	19.00
<b>Ne</b>	10	Neon	20.18
<b>Ar</b>	18	Argon	39.95
<b>Kr</b>	36	Krypton	83.80
<b>Xe</b>	54	Xenon	131.30
<b>Rn</b>	86	Radon	(222)

Transition elements: valence electrons not in outer shell

<b>Zn</b>	30	Zinc	65.38
<b>Cu</b>	29	Copper	63.55
<b>Ni</b>	28	Nickel	58.70
<b>Co</b>	27	Cobalt	58.93
<b>Fe</b>	26	Iron	55.85
<b>Mn</b>	25	Manganese	54.94
<b>Cr</b>	24	Chromium	52.00
<b>V</b>	23	Vanadium	50.94
<b>Sc</b>	21	Scandium	44.96
<b>Ti</b>	22	Titanium	47.90
<b>Zr</b>	40	Zirconium	91.22
<b>Hf</b>	72	Hafnium	178.49
<b>Rf</b>	104	Rutherfordium	261.11
<b>Y</b>	39	Yttrium	88.91
<b>La</b>	*	(see below)	
<b>Ce</b>	58	Cerium	140.12
<b>Pr</b>	59	Praseodymium	140.91
<b>Nd</b>	60	Neodymium	144.24
<b>Pm</b>	61	Promethium	(145)
<b>Sm</b>	62	Samarium	150.4
<b>Eu</b>	63	Europium	151.96
<b>Gd</b>	64	Gadolinium	157.25
<b>Tb</b>	65	Terbium	158.93
<b>Dy</b>	66	Dysprosium	162.50
<b>Ho</b>	67	Holmium	164.93
<b>Er</b>	68	Erbium	167.26
<b>Tm</b>	69	Thulium	168.93
<b>Lu</b>	71	Lutetium	174.97
<b>Yb</b>	70	Ytterbium	173.04
<b>La</b>	*57	Lanthanum	138.91
<b>Ce</b>	58	Cerium	140.12
<b>Pr</b>	59	Praseodymium	140.91
<b>Nd</b>	60	Neodymium	144.24
<b>Pm</b>	61	Promethium	(145)
<b>Sm</b>	62	Samarium	150.4
<b>Eu</b>	63	Europium	151.96
<b>Gd</b>	64	Gadolinium	157.25
<b>Tb</b>	65	Terbium	158.93
<b>Dy</b>	66	Dysprosium	162.50
<b>Ho</b>	67	Holmium	164.93
<b>Er</b>	68	Erbium	167.26
<b>Tm</b>	69	Thulium	168.93
<b>Lu</b>	71	Lutetium	174.97

IA IIA IIIB IVB VB VIB VIIB VIIIB IIB IIIA IVA VA VIA VIIA VIIIA

Lanthanide (Rare Earth) Elements

<b>La</b>	*57	Lanthanum	138.91
<b>Ce</b>	58	Cerium	140.12
<b>Pr</b>	59	Praseodymium	140.91
<b>Nd</b>	60	Neodymium	144.24
<b>Pm</b>	61	Promethium	(145)
<b>Sm</b>	62	Samarium	150.4
<b>Eu</b>	63	Europium	151.96
<b>Gd</b>	64	Gadolinium	157.25
<b>Tb</b>	65	Terbium	158.93
<b>Dy</b>	66	Dysprosium	162.50
<b>Ho</b>	67	Holmium	164.93
<b>Er</b>	68	Erbium	167.26
<b>Tm</b>	69	Thulium	168.93
<b>Lu</b>	71	Lutetium	174.97

Actinide Elements

<b>Ac</b>	*89	Actinium	227.03
<b>Th</b>	90	Thorium	232.04
<b>Pa</b>	91	Protactinium	231.04
<b>U</b>	92	Uranium	238.03
<b>Np</b>	93	Neptunium	237.05
<b>Pu</b>	94	Plutonium	(244)
<b>Am</b>	95	Americium	(243)
<b>Cm</b>	96	Curium	(247)
<b>Bk</b>	97	Berkelium	(247)
<b>Cf</b>	98	Californium	(251)
<b>Es</b>	99	Einsteinium	(252)
<b>Fm</b>	100	Fermium	(257)
<b>Md</b>	101	Mendelevium	(258)
<b>No</b>	102	Nobelium	(259)
<b>Lr</b>	103	Lawrencium	(260)